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MICROBIAL AND ENZYME ACTIVITIES OF SALINE AND SODIC SOILS

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ABSTRACT

Salinization and sodication are abiotic soil factors, important hazards to soil fertility and consequently affect the crop production. Soil salinization is of great concern for irrigated agriculture in arid and semi-arid regions of the world; sodicity is characterized by an excessively high concentration of sodium (Na) in their cation exchange system. In recent times, attention has been turned to study the impacts of these factors (salinity and sodicity) on soil microbial activities. Microbial activities play central role in degradation and decomposition of soil organic matter, mineralization of nutrients and stabilization of soil aggregates. To understand the ecology of soil system, therefore, it is important to be conversant with the soil microbial activities, which show quick response to little change in the soil environment. Microbial activities (generally measured as C–N dynamics, soil respiration–basal respiration, or CO_2 emission), microbial abundance, microbial biomass, quotients (microbial and metabolic) and microbial community structure, and soil enzymes have been considered as potential indicators to assess the severity of the land degradation and the effectiveness of land use management. Therefore, it is important to synthesize the available information regarding microbial activities in use and management of salt-affected soils. The reclamation and management of such soils and their physico-chemical properties have been reviewed well in the literature. In this review, an attempt has been made to compile the current knowledge about the effects of soil salinization and sodication on microbial and enzyme activities and identify research gaps for future research. Copyright © 2015 John Wiley & Sons, Ltd.

KEY WORDS: exchangeable sodium percentage; microorganisms; microbial activities; salt-affected soils; soil enzymes

INTRODUCTION

Deterioration of soil quality by soil salinization and sodication is a global problem. Salt affected soils, are commonly occurring in arid and semi-arid regions (Pandey et al., 2011; Ferreira et al., 2015) and gradually spreading in various regions of the world (Lambers, 2003). Saltaffected soils occupy approximately 3% of the world's total geographical area; 402 and 434 million hectare (Mha) lands are classified as saline and sodic, respectively. There are 137 Mha (5% of the total geographical area) sodic soils in the Asia-Pacific region (Shukla et al., 2011). Furthermore, extensive agricultural practices with poor irrigation and drainage practices may lead to manifold increase in this area in the forthcoming decades (Lambers, 2003). Salinization and sodication affect physico-chemical, biological, and biochemical properties of soil (Gill et al., 2008; Rietz & Haynes, 2003; Singh et al., 2013b, 2013c), and affecting the crop productivity to a significant extent (Gill et al., 2009). Despite the extensive studies on the effects of salinity and sodicity on soil properties (Qadir et al., 2007) and crop growth, little is known about their possible effects on soil microbial activities. Soil microbial activities (SMAs) are considered as key players in the formation of soil structure (Elliot, 1986; García-Orenes et al., 2010), biodegradation of organic matter (Quails & Raines, 1992), nutrient cycling,

and energy transformation (Coleman et al., 1985); therefore, their understanding in salt-affected soils is of paramount importance. The effects of fluctuations in electrical conductivity (EC; salt solutions: salinization) and exchangeable sodium percentage (ESP) or sodium adsorption ratio (SAR) (sodication) on microbial activities have been examined by several workers. Both positive and negative effects were observed in these studies with respect to salinity and sodicity, and therefore, reconciliation of proper mechanisms is yet to be resolved. The reduced and/or altered microbial activities in saline and sodic soils are likely either due to direct toxic effects of salts on microbial communities (Rietz & Haynes, 2003) or the salinity and sodicity may exert toxic effects on vegetation (crop, grass, and tree) (Wong et al., 2010), which in turn, leading to decreased organic matter inputs (crop residue, litter, and fine roots) in the soil and, consequently, significant decrease in microbial activities (Singh et al., 2015).

Two contrasting procedures are found in saline and sodic soils that affect soil properties (Sumner, 1993; Sumner *et al.*, 1998; Wong *et al.*, 2008). First is decreasing water availability (less osmotic potential) of the soil solution with increasing salinity and second is increasing mineralization of organic matter (carbon loss) with increasing sodicity (Rengasamy, 2006a, 2008; Wong *et al.*, 2008, 2010). Changes in soil chemical properties and osmotic potentials due to fluctuations in salinity are potential reasons affecting the microbial growth and activities (Wong *et al.*, 2010; Setia *et al.*, 2011a; Guénon *et al.*, 2013). The most notable is exosmosis or dehydration (Rengasamy, 2006b, 2008).

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nisms of remediation of salt stress (Sumner, 1993, 1995; Qadir *et al.*, 2007; Dendooven *et al.*, 2010; Wong *et al.*, 2010), are available, but none has included microbial and enzyme activities. The objective is to review the literature, provide facts, concepts, and mechanisms of microbial and enzyme activities in salt-affected soils as well as to find out knowledge gaps for directing the future researches to understand the microbiology of salt-affected soils.

However, the accumulation of osmolytes in microbial cells

enhances salt tolerance ability of microbes (Hagemann,

2011; Asghar et al., 2012) to combat with low osmotic

and metric potentials in the soil solution. On the other hand,

in sodic soils, the available soil organic matter (SOM) is free

from physical protection (poor aggregation) and undergoes

for higher biological mineralization (Abiven et al., 2009;

Wiesmeier *et al.*, 2012). Higher accessibility (higher mineralization) of SOM for microbial decay in sodic soil may fur-

ther deteriorates the fertility and physical stability of the soil

(Nelson & Oades, 1998; Wong et al., 2010; Fterich et al.,

2014). At high sodicity, availability of organic matter and

metabolic energy, required for microbial assimilation, are re-

duced consequently; microbial cell lyses takes place and soil

effects of salinization and sodication on soil physico-

chemical properties, carbon-nitrogen dynamics, and mecha-

Indeed, several comprehensive reviews, dealing with the

becomes poor in microbial activities.

SALINE AND SODIC SOILS

Soils with high concentration of salts (EC), high pH, ESP, and SAR are collectively called as salt-affected soils (Soil Salinity Staff, 1954; Qadir et al., 2007). In general, first soil salinization takes place, and then saline soils become sodic as Na replaces Ca and Mg from the cation exchange sites that causes dispersion of clay particles (Sumner, 1993) and, thereby, increases oxygen stress in the soils (Richards, 1954; Sumner, 1993; Naidu & Rengasamy, 1995). Sodic soils contain high concentrations of sodium carbonate (Na₂CO₃), sodium bicarbonate (NaHCO₃), and sodium chloride (Abrol & Bhumbla, 1971). Sodic soils generally have poor physical properties, with low water permeability (Qadir et al., 2007). Sodicity has substantial effects on soil physical properties, causing deterioration of soil structure because of change in slaking and dispersion behavior of clay minerals upon wetting and swelling, crusting, and hardsetting on drying (Sumner, 1993; Qadir et al., 2007; Wong, 2007; Wang et al., 2015). This affects permeability, infiltration, and hydraulic conductivity of soils. These soils can develop naturally from saline soils, with their development related to the underlying parent material, climatic change, or as a result of leaching of salts from a saline soil (Wong, 2007). Sodic soils do not sustain significant vegetation cover (Singh et al., 2013a), therefore, receive less organic matter inputs and consequently have lower organic carbon (Wong et al., 2008). Low organic-matter input and high Na stress enhance the mineralization of meager quantity of inherent organic matter (Nelson *et al.*, 1998) and induce aggregate destruction (Abiven *et al.*, 2009). Sodic soils with poor aggregate stability provide inadequate protection to the SOM. Readers interested in details of clay behavior (swelling, slaking, dispersion, and flocculation); osmotic potentials; and levels of pH, EC, ESP, and SAR in saline and sodic soils are referred to other reviews (Richards, 1954; Sumner, 1993; Naidu & Rengasamy, 1995; Nelson *et al.*, 1998; Qadir *et al.*, 2007).

Saline soils with high concentrations of salts that had its adverse effect on nutrient availability causing a pseudonutrient deficiency in the soil water to crops and trees. Saline soils showed good physical properties having fine particles bound with organic matter and form macroaggregates. Furthermore, this causes flocculation that is beneficial for crop root growth (Nelson *et al.*, 1997; Nelson & Oades, 1998; Wong, 2007). Rietz & Haynes (2003) stated that soils with high electrolyte concentration are categorized as saline, but these are physically stable with greater aggregate stability, water infiltration rate, clay flocculation, and less density. Saline soils under agricultural practices receive organic inputs as crop residues, but poor quality of irrigation water and extensive chemical fertilization lead to salinization.

MICROBIAL ACTIVITIES IN SALINE AND SODIC SOILS

Microbial activities can be determined by decomposition of SOM (Rietz & Haynes, 2003), release (mineralization) of nutrients (N, P, and S) (Nelson et al., 1998), carbon dioxide (CO₂) emission, and transfer of C from SOM to microbial biomass (Wong et al., 2010; Balota et al., 2013; Campos et al., 2013). The important parameters used to explain soil microbial activities under various stresses are describes in Table I. These activities are dependent on many factors including both biotic and abiotic (Yuan et al., 2007; Egamberdieva et al., 2009). In general, these activities are likely to be optimum in moist neutral (pH around 7) soils at high temperature with adequate organic matter in comparison with cool, dry, and alkaline or acidic soils (Rukshana et al., 2012). The increased salt content (saline soils) and decreased structural stability (sodic soils) along with other chemical alterations affect soil quality in many different ways (Figure 1). Therefore, it would be interesting to relate microbial activities with changes in salinity (EC and water content), sodicity (ESP, SAR, pH, and clay mineralogy), and quality and quantity of SOM.

Soil Respiration (CO_2 emission)/ Carbon Mineralization in Saline and Sodic Soils

The SOM decomposition and the magnitude of carbon dioxide (CO₂) emission [soil respiration (SR)] from soil are affected by salinity and sodicity (Wong *et al.*, 2008; Dendooven *et al.*, 2010). However, this might depend on the quality and quantity of organic matter (Nelson & Oades,

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Table I. Description of microbial activities used to elicit the effects of salinity and sodicity

| Microbial activity | Description | Reference |
|---|--|--|
| Soil respiration (CO ₂ emission)/ basal soil respiration | Soil respiration (SR) refers to the emission of carbon dioxide (CO_2) during respiration soil organisms (microbes and fauna) as well as plant roots. SR, as a key factor influencing the dynamic change of CO_2 concentration in atmosphere, is the biggest carbon efflux process after photosynthesis in the terrestrial ecosystem. | Davidson <i>et al.</i> (2006), & Wang & Fang (2009) |
| Soil microbial biomass (SMB) | Soil microbial biomass is a measure of the mass of the living component of soil organic matter. This SMB is responsible for decomposition of plant and animal residues and soil organic matter to release CO2 and plant available nutrients. | Anderson & Domsch (2010) |
| Microbial quotient (C _{mic} /C _{org}) | The microbial quotient is the ratio of the SMB-C to SOC (C_{mic}/C_{org}) and indicates the ratio of the living fraction of SOC relative to the non-living fraction. | Anderson & Domsch (1993) |
| Metabolic quotient (qCO ₂) | The metabolic quotient (q CO ₂), also known as the respiratory quotient, is frequently used to determine stress in the microbial population (Anderson & Domsch, 1993) and measures the ratio of soil respiration to the SMB. | Anderson & Domsch (1993) |
| Soil enzyme activities | Soil enzymes play key biochemical processes in the soil system to control all processes involved in decomposition of organic matter and nutrient cycling. They have been recognized as proximate agents for mineralization of organic matter. Different intracellular and extracellular enzymatic components contribute to the overall | Sinsabaugh et al. (2008), Yao et al. (2009) Gianfreda & Rao, 2011; Nannipieri et al., 2012 |



Figure 1. 1: Salinity and sodicity may exist together in soil as sodicity represent combined effects of salinity (EC) and sodicity (ESP). 2: Increase in salinity has its osmotic influences. 3: Osmotic influences cause microbial cell lyses or tolerant microbes can accumulate cell solute as osmolytes. 4: Altered microbial community affects decomposition of organic matter. 5: Addition of organic residues can stimulate the microbial biomass and activities and lighten salinity stress. 6: Increase in salinity enhances CO_2 emission because fewer microbes respire more and decompose readily available organic molecules under salinity stress. 7: CO_2 produced by microorganisms may be dissolved in the soil solution, thus underestimating the actual CO_2 evolution. 8: Dissolution of $CaCO_3$, available in sodic soils in excess, form CO_2 , thereby leading to an overestimation of CO_2 emission. 9: Salinity affects N-mineralization through its direct effect on bacteria that produces ammonia from organic matter. 10: Increase in microbial biomass due to addition of organic residues (at 5) may immobilize inorganic N and reduce the availability of soil inorganic N. 11: Ammonia may be lost through volatilization. 12: There may be an increase in inorganic N due to fixation of atmospheric N. 13: Sodicity affects soil structure (size, shape, and arrangements of particles and pores) and behavior (slaking, dispersion and swelling). 14: This structural change in sodic soil facilitates the accessibility of organic matter for microbial decomposition. 15: Soil organic matter can affect the sodicity, and sodicity also influence soil organic matter. EC, electrical conductivity; ESP, exchangeable sodium percentage. This figure is available in colour online at wileyonlinelibrary.com/journal/ldr

1998), level of salinity and sodicity in soil (Pathak & Rao, 1998; Rietz & Haynes, 2003; Tejada *et al.*, 2006), type of salt present or added during incubation study (Li *et al.*,

2012), interaction of organic matter and microorganisms with clay mineral surface (Nelson *et al.*, 1997), and soil structure.

CO₂ Emission or C Mineralization in Saline Soils

Soil salinity is developed by a relatively high electrolyte concentration of different salts, especially, NaCl and CaCl₂, in soil solution. An increase in salt concentration results in less available water for microbial and plant growth. Salinity, beyond tolerance of microbial fauna and plants, causes death of microbial cells and degradation of plant tissues because of dehydration, toxicity to soil enzymes, and inhibits other soil processes. The reduction in microbial population and organic inputs affect decomposition process and availability of essential plant nutrients and, thereby, plant productivity. Simultaneously, under salinity stress, microbial fauna try to acclimatize with changing environment and take energy from easily decomposable organic substrates that are produced because of death of microbial cells and plant tissues, and other available labile organic molecules. These microbes increase mineralization of organic substrates, evolve more CO₂, and get energy for their own sustenance. This shift in microbial activities is up to limited level of salinity as an adaptation mechanism beyond that salinity would always be detrimental to microbial activities. Fluctuations in salinity in field conditions depend on irrigation water quality, irrigation frequency, nature of crop (deep rooted or shallow rooted) being cultivated, duration of rainy and hot seasons, field application of organic and chemical fertilizers, and other abiotic factors. Most of the research carried out to know the impact of salinity on microbial activities (C and N mineralization) was based on incubation studies. In these studies, different salts in various concentrations and organic matter (residues or compost or litter), either alone or in combination, were applied to non-saline soils. This artificially created salinity may or may not be an actual factor for changes in microbial activities especially at early stages of incubation study and may not be that much similar to natural field conditions (Macci et al., 2013).

Factors affecting CO₂ emission from saline and salinesodic soils have been correlated with soil properties including, osmotic potential and sodium adsorption ration (Setia et al., 2010; Setia et al., 2011b; Setia et al., 2011c), soil texture (Chowdhury et al., 2011c; Mavi et al., 2012) and water content (Setia et al., 2011a), soil organic carbon stocks (Setia et al., 2012a), and residue composition and water potential (Setia et al., 2012b; Setia & Marschner, 2013a; Setia & Marschner, 2013b). Several other researchers have also investigated the effect of salinity and sodicity on SR (CO₂ evolution) in laboratory (Pathak & Rao, 1998; Beltrán-Hernández et al., 2007; Wong et al., 2008, 2009) and field studies (Chander et al., 1994; Rietz & Haynes, 2003; Yuan et al., 2007; Singh et al., 2012a). In all these studies, when salinity and sodicity have been induced, results were rather contradictory; both increase (Nelson et al., 1996; Wong et al., 2008, 2010) and decrease (Pathak & Rao, 1998) in SR have been reported. It seems that SR (CO₂ emission) besides salinity and sodicity depends on integrated effects of anthropogenic and environmental factors in a complex phenomenon.

An increase in salinity up to an EC of 5 dS m⁻¹ in the silty loam soil either with NaCl or CaCl₂ has no significant differences between two salts on SR (Setia *et al.*, 2010). Earlier study (Agarwal *et al.*, 1971) indicates that Ca⁺⁺ enhances the microbial metabolic activities, while from the later study, (Setia *et al.*, 2010) it seems that Na⁺ and Ca⁺⁺ did not differ in their toxicity to SMAs. Because SR is the result of CO₂ produced from decomposition of SOM due to microbial action and root respiration, we may not differentiate the impacts individually. Setia *et al.* (2010) argued that values of CO₂ released are not always accurate as some of CO₂ may be retained in the soil water and some may come as a result of SOM decomposition. Thus, there is respective chance of underestimation as well as overestimation CO₂ evolution.

Yuan et al. (2007) determined lower CO2 emission in saline soils with highest organic carbon content in comparison with non-saline soils. This may be explained with two ways. First, organic matter is less available for microbial decomposition because of good aggregation in saline soils. Second, if this value of CO_2 is similar to other saline soils in which soil organic carbon (SOC) is less, then certainly, the salinity stress affects CO₂ emission. Setia & Marschner (2013a) reported inconsistent response in CO₂ emission in saline soils with effect of wheat and pea residues; they postulate that, in saline soils, type of residue did not affect SR, and in nonsaline soil, the decrease in respiration with increasing salinity was affected by type of residue. The influence of residue on SR may be diluted under stress conditions irrespective of the quality of substrate. However, in case of non-saline soils, residue may affect the SR as no other stress will be able to dilute its affect. This may also depend on type, quantity, chemistry, and moisture content of added residue. Moist residue with less C: N ratio can decompose easily than dry with high C:N ratio. The concentration and type of organic carbon (particulate organic carbon and humus-carbon) and clay minerals are also influenced the CO₂ emission to different magnitudes whether soil is saline or not (Setia et al., 2011b). The high salt concentration and high water in saline and non-saline soils will influence microbial activities differently. If the water content are relatively higher in saline soils, the high salt concentration strongly reduces microbial activities (Setia & Marschner, 2013b). The amendment of such soils with organic residue enhances their subsequent mineralization with a related increase in CO₂ emission and soil aeration (Muhammad et al., 2007) probably because of stimulus in their enzymatic activities (Lakhdar et al., 2009). The effect of different amounts of mature-wheat residue on SR in saline environment along with addition of NaCl was investigated by Setia et al. (2010). SR was influenced at all residue addition rates; however, it was more pronounced at low residue rates. The small increase in microbial biomass at low residue addition rates was unable to decompose the residues at high rates and/or the low density of residues in the soil resulted in spatial separation of microbes and substrates (Setia et al., 2010).

The CO_2 flux was lowest in saline soils in comparison with peat, non-calcareous, terrace, and acidic sulfate soils after addition of wheat, mungbean, and poultry manure. Loss of carbon in saline soil was 30% and 39% after addition of wheat and mungbean residues, respectively (Khalil et al., 2005). This indicates that saline soil affects C mineralization to various degrees after addition of different organic residues and salts. This also changes soil pH (Rukshana et al., 2012, 2013) in different ways, directions, and magnitudes depending on initial soil pH, chemical composition of organic residue and microbial activities especially CO₂ emission (Khalil et al., 2001). The rate of C mineralization in the saline soils may decrease or increase according to change in soil pH and, thereby, change the other soil properties (Khalil et al., 2005). It appears that organic matter with high C: N might result in a reduced C loss because of slow decomposition rate under aerobic saline conditions; this further depends on the quality and quantity of C source in the soil and extent of C released (Khalil et al., 2001). For instance, CO_2 was higher with $5 g C kg^{-1}$ soil than with 2.5 gC kg^{-1} soil (quantity), and glucose treatment enhanced respiration over the cellulose treatment (quality). The microbes in saline conditions decompose cellulose more easily than glucose (Elmajdoub & Marschner, 2013). The effect of salinity on potential of microbes to mineralize C from various sources and states is also not well known. In this context, the response of SMB and activity was determined in an incubation experiment to salinity when supplied with two different carbon forms (glucose and cellulose) (Elmajdoub & Marschner, 2013). Yan & Marschner (2012) reported that microbes isolated from saline soils showed less tolerance to increasing salinity than those isolated from non-saline soils. It is assumed that addition of residue may increase MB because of a small set of microbes in all soils feed on added residues in saline soils (Yan & Marschner, 2012). They further extended their work (Yan & Marschner, 2013) and established an experiment with a range of saline levels created artificially in which pea straw was added as a nutrient source. At artificially created salinity levels, respiration recorded similar to that of soils that had originally lower salinity. This suggests that structure and function of soil microbial community can rapidly increase when the substrate is added and salinity decreased.

Despite the fact that salinity changes with water content in soil, little is known about effects of moisture content on SR and microbial activities. Fluctuations in water content (water potential) tremendously affect SMAs (Setia & Marschner, 2013a; Yan & Marschner, 2013). In soils with high concentration of salts may show more pronounced effects. Wetting of dry soils always induce gaseous emissions including CO₂, which may decrease with increase in time. This trend in saline soils is thought to be due to dispersion of clay minerals, exposure of organic substrates, release of osmolytes due to cell lysis, and breakdown of aggregates (macro to micro) (Chowdhury et al., 2011b). Several other factors may also affect the SMAs like interval between rewetting and drying, initial osmotic potential, soil fertility, structure, addition of substrate, and irrigation levels. An increase in dryingrewetting cycles lowers flush of CO2 (Chowdhury et al., 2011b). The magnitude of alteration in aggregates because of introducing water in soil affects microbial activities. The sudden wetting of soil destroys aggregates and exposes organic matter for microbial assimilation, but at later stages, wetting does not affect aggregation to the similar scale (Chowdhury *et al.*, 2011a).

Setia & Marschner (2013b) studied the effect of osmotic, matric, and water potential, and substrate composition on microbial activity and biomass in saline soils amended with wheat and pea residues. In this study, cumulative respiration decreased with decreasing water potential, and the relative decrease in microbial activity suggests that the small biomass in saline soils does not tolerate low osmotic potential (Setia & Marschner, 2013a). Hence, drying of soil would have negative effect on microbial activity in saline than in non-saline soils (Setia & Marschner, 2013b). Similarly, Mavi & Marschner (2012) studied the impact of multiple drying and wetting on microbial respiration, dissolved organic carbon, and microbial biomass in salt-affected soils. Rewetting of dry soils resumed respiration in the slightly saline soils but could not in moderately saline soils, which suggest the better substrate utilization by microbes in less saline soils. Iwai et al. (2012) investigated the effect of alternative wetdry tropical climatic conditions. They found an increase in exchangeable Na⁺ during dry season that indicates upward moment of salts on the soil surface during the dry season. The significant negative exponential relationships between EC and microbial activities were observed during dry and wet seasons.

CO2 Emission or C Mineralization in Sodic Soils

Sodic soils are characterized with a very low SOM because of less C sequestration and high decomposition rates. Sodicity may influence the decomposition of SOM both directly and indirectly to have a relatively low SOM in the soil. Sodicity influence the structure (shape, size, and arrangements of particles and pores in soil) and, thereby, minerals' behavior (slaking, swelling, and dispersion), which plays an important role in microbial and faunal activities and trophic interactions (Nelson & Oades, 1998). A high level of sodium (Na) makes SOM more accessible for microbial decomposition by dissolving and dispersing. The positive effect of Na⁺ on decomposition could be greatest for small anionic substrates and least for particulate uncharged substrates. The anaerobic condition caused by sodicity may decrease mineralization in the field through their negative effect on microbial activities. As the level of sodicity determines the quality and quantity of SOM, the reciprocal effects of SOM on soil sodicity are also noticed in different ways. The cation exchange capacity and the selectivity of exchange sites for Ca over Na, and ability of SOM to donate protons have significant effects on sodicity. In conclusion, sodicity may increase the mineralization of SOM by making it more accessible for microbial decomposition, but at the same time, it may also decrease mineralization in submerged conditions because of the effects of anaerobic condition on microbial community (Wu et al., 2013).

Amending sodic soils with Ca salts decreased C mineralization (less CO_2 emission) under both laboratory and field conditions (Muneer & Oades, 1989a, 1989b) because Na⁺ is more effective in prevailing dispersion of clay minerals than Ca⁺⁺ (Sumner, 1993). Decrease in SR might also be due to decreased dispersion of SOM and clay particles by amending soils with Ca⁺⁺. However, addition of organic matter facilitates the dissolution of CaCO₃; in that case, CO₂ evolution is increased. The evolved CO₂ may also be trapped in the soil solution through formation of Ca (HCO₃)₂ (Oren & Steinberger, 2008; Macci *et al.*, 2013). There is a need to relate the CO₂ evolution with total (insoluble) and active (soluble) CaCO₃ in the soil with different salinity and sodicity levels.

N Mineralization in Salinity and Sodicity

The mineralization of N in salt-affected soils is a subject of deep insight because of preliminary controversial studies (Laura, 1974, 1975, 1976; Nelson et al., 1996; Pathak & Rao, 1998). Salinity and sodicity of soil have its effects on C and N mineralization, which are crucial for plant growth and soil fertility. Ammonification is stimulated at low salt concentration and inhibited at higher amounts of salts, but nitrification is very sensitive because ammonifiers are more tolerant than nitrifiers (Westerman & Tucker, 1974; McCormick & Wolf, 1980; McClung & Frankenberger, 1987; Pathak & Rao, 1998). Laura (1974) reported that ammonification is continued to increase with increasing salinity. However, the ammonification apparently increased up to EC $70 dS m^{-1}$, while nitrification was stimulated only up to EC $16 dS m^{-1}$ and completely inhibited at EC $26 dS m^{-1}$ (Pathak & Rao, 1998). Therefore, salinity causes a significant decrease in N mineralization. Stimulation of ammonification at low salinity may be explained on the basis of solubilization of SOM and osmotic effects on microbes (Broadbent & Nakashima, 1971) releasing carbonaceous substrates (Pathak & Rao, 1998). Increased solubilization of SOM at high pH may lead to increase N mineralization. In sodic soils, NH₄-N was not detected, and there was significant increase in NO3-N with increasing soil pH and ESP (Pathak & Rao, 1998). This reveals that increasing sodicity stimulated N mineralization but the values of net N mineralization, as a fraction of organic N, showed a small decrease with increase in soil pH (Pathak & Rao, 1998). For instance, increase in NO₃-N concentration with increase in pH or ESP is not real stimulation of nitrification. Rietz & Haynes (2003) presented negative relationship of N mineralization and ammonification with EC, SAR, and ESP. This indicates that increase in salinity and sodicity is harmful to the most of the microbes, and in that case, less metabolically efficient soil microbes mediate soil processes including C, N, P, and S mineralization. The decomposition of SOM is generally reduced with effect of salinity, but it may be unaffected or increased with effect of sodicity (Nelson et al., 1996). In any case, if the decomposition rate is decreased by salinity or sodicity; it is directly associated with the inhibitory microbial growth and activity (Rietz & Haynes, 2003).

The concentration of NO₃–N increased (stimulation of nitrification) in sludge amended soil with high salt content and pH 10·2 (Dendooven *et al.*, 2010), while high salt contents and pH are known to inhibit oxidation of NH⁴₄ and NO²₂ (Oren, 1999, 2002). The amount of NH⁴₄ and easily decomposable organic material available in soil amendment also affect the rate of nitrification and ammonification in saline and sodic soils (Dendooven *et al.*, 2010). The decrease in NH₄–N might be due to NH₃ volatilization at high pH (Sommer & Ersbøll, 1996) and adsorption of NH⁴₄ on soil particles (Drury *et al.*, 1991). In other cases, an increase in NH₄–N with increasing salinity and sodicity appears that microbes in these soils immobilized NH⁴₄–N in excess of what was necessary for metabolic activity (Dendooven *et al.*, 2010).

Soil Microbial Biomass, Population, and Quotients with Salinity and Sodicity

It is clear from previous sections that, when microbial population is under stress, more C is lost through respiration rather than being converted to their own growth and humus formation. In most of cases, the mineralized N is decreased with increasing salinity and sodicity levels in soil. It would be interesting to know how microbial biomass, population, and potential enzyme activities are responding to salinity and sodicity stress. The quality and quantity of organic residue (substrates) and soil environment play important roles in microbial assimilation of soil organic carbon. Low SOM and high salinity may establish an undesirable environment for the development of bacterial and fungal community, and in contrast, high organic matter and low salinity may favor the fungal growth (Yuan et al., 2007). Muhammad et al. (2006) found a higher fraction of fungal biomass in the more saline soil compared with less saline soils. Besides, in tropical and subtropical climate where saline and sodic soils are found, decomposition process is largely carried out by bacteria; whereas in temperate climate, fungi may perform major role in this process.

Salinity and sodicity exert detrimental effects on both microbial biomass and thereby on metabolic quotient. Therefore, the metabolic quotient may indicate the response of microbial population under stress (salinization and sodication) (Wong et al., 2008, 2005). Chander et al. (1994) reported that MBC reduced to one-third in the soils treated with 16 mel^{-1} residual sodium carbonate (RSC) irrigation water in comparison with that of treated with 2.8 mel^{-1} RSC irrigation water. This study indicates that an increase in sodicity over a critical level affects microbial biomass identically. Yuan et al. (2007) reported significant negative exponential relationship between EC (salinity) and microbial biomass C and N and microbial quotients. The highest values of qCO_2 in the most of the saline and sodic soils indicate low substrate availability and poor functioning of microorganisms (Anderson & Domsch, 1990; Wong et al., 2009) and variations in the composition of the microbial communities (Tripathi et al., 2006). The low percentage of MBC to that of SOC demonstrates that salinity

is detrimental to microorganisms (Yuan *et al.*, 2007); however, it is also noticed that considerable microbial activity persisted in saline soils because of accumulation of osmolytes in salt tolerant bacteria (Tejada *et al.*, 2006; Yuan *et al.*, 2007; Chowdhury *et al.*, 2011c; Asghar *et al.*, 2012). Chowdhury *et al.* (2011) and Setia *et al.* (2011a, 2011b) reported the decrease in SR with increasing salinity even after addition of substrate. It was speculated that higher metabolic quotient and low MBC at high soil salinity might be due to either a shift in microbial community with less catabolic activity or reduced efficiency to utilize substrate.

With the increasing level of salinity a consistent decrease in SR was reported, while MBC was unaffected (Ghollarata & Raiesi, 2007). Wong et al. (2008) also observed highest SR in low level of soil salinity and lowest in high level of soil salinity, while SMB was highest in the high-salinity soils and lowest in low-salinity soils. They postulated that highest microbial biomass in high-salinity soils was due to more substrate availability, high salt concentrations, high dispersion of aggregates, and high solubility of organic matter. Thus, reduced efficiency of substrate utilization by the soil microflora could be a mechanism by which salinity depresses microbial respiration even with added organic matter. Singh et al. (2012a), in their field study, reported lowest microbial biomass and SR (CO2 efflux) in sodic soils with 95% (ESP); values of SMB and SR increased with decreasing sodicity when the soil was reclaimed substantially by afforestation and cropping over the years.

Osmotic desiccation in saline environment may lead to microbial cell lysis of less tolerant microbes (Dinesh *et al.*, 1995). Salinity and low nitrogen content may limit microbial activity and residue decomposition rate (Elgharably & Marschner, 2011). Elgharably & Marschner (2011) studied the effects of NH₄–N or NO₃–N and addition of lupin residues on microbial biomass and activity in a saline sandy loam. Soil microbial biomass (SMB) decreased with increasing salinity but significantly increased with residue addition. The greater C availability in the lupin residue amended saline sandy loam soil and stimulated microbial activity and

biomass with greater N demand, thus promoted immobilization of NO₃. Rietz & Haynes (2003), in agreement with Galinski (1995) and Oren (1999), also proved that, in saline soils, osmotic stress limits the microbial growth and activity and causes dehydration of microorganisms. In sodic environment, however, Na toxicities and high pH may inhibit microbial growth (Rietz & Haynes, 2003). Increase in SMB and SR was observed with addition of two organic wastes (cotton gin crushed compost, CGCC, and poultry manure, PM) to saline soils by Tejada et al. (2006). However, SMB and SR were higher in PM than in CGCCamended soils, while decrease in ESP was higher in CGCC-amended soils. It was postulated that this might be because of greater labile carbon, N, and P content in PM than in CGCC. This might also be attributed to (i) microbial growth and enzyme activities exist in these organic wastes, (ii) higher sodicity (ESP) and substrate availability in PM than CCGC-amended soils, and (iii) dispersion of soil aggregates (Wong et al., 2008).

I have derived a generalized pattern of CO₂ emission, N-mineralization, microbial biomass, and enzyme activities (Figure 2) through reviewing the data from the literature cited in preceding sections. It is observed that microbial activities may differ in response to substrate in a particular soil condition; when we add the substrate, it may show entirely different response. In the absence of adequate organic matter (organic substrate), CO₂ emission increases with increasing salinity and sodicity, while N-mineralization, microbial biomass, and enzyme activities decreased (Figure 2a). On the other hand, the case may be vice versa when an organic substrate (ranging from organic molecules to crop residues or organic matter) is added externally to salt-affected soils CO₂ emission increases with respect to increasing salinity and sodicity and N-mineralization, microbial biomass and enzyme activities increase to attain an equilibrium (steady state). Therefore, it is clear that response of microbial activities to soil salinization and sodication depends on quality and quantity of inherent organic substrates as well as externally added to salt-affected soils.



Figure 2. The effect of salinity (EC) and sodicity (ESP or SAR) on microbial activites [A] without substrate addition and [B] after substrate addition. EC, electrical conductivity; ESP, exchangeable sodium percentage. This figure is available in colour online at wileyonlinelibrary.com/journal/ldr

Microbial Community Structure and Function with Salinity and Sodicity

Identification of microorganisms through the traditional culture dependent techniques remains limited to less than 1% of the bacterial diversity. However, several physiological (community level physiological profiling), biochemical (phospholipid fatty acid analysis; PLFA) (Singh et al., 2004), and molecular (microbial DNA fingerprinting) methods (Dubey et al., 2006) (culture-independent techniques) are in use to investigate the soil microbial diversity (Singh et al., 2004). Rousk et al. (2011) studied the influence of soil salinity on soil bacterial communities in a range of salt levels. The clear concentration response relationship was established between bacterial growth and soil salinity, demonstrating an accurate assessment of bacterial tolerance. They reported that soil salinity was not a decisive factor for bacterial growth and thus for structuring the decomposer community in the studied soils. Similarly, Asghar et al. (2012) reported that microbial community composition (based on PLFA) and activities of microorganisms isolated from saline and non-saline soils respond similarly to changes in salinity. This might be because of modified physiological and structural characteristics of the microbial communities in saline environment (Zahran, 1997: Ventosa et al., 1998; Rietz & Haynes, 2003; Rao & Pathak, 1996). However, these findings did not consider other factors existing in natural field soils.

Wu et al. (2012) ameliorated seashore saline soil by planting castor bean (Ricinus communis L.) for two growing seasons. After 2 years of growth, soil salinity was 0.92% lower than that unplanted soil; the activity and biodiversity of soil microbial communities (based on community level physiological profiling) and functional and metabolic diversity of halophilic, phosphate-solubilizing, potassium-solubilizing, cellulose decomposing, ammonifying, and nitrogen-fixing bacteria increased markedly in the planted plots in comparison with control plots. The increase in functional microorganism that can fix atmospheric nitrogen, solubilize phosphate, potassium, or decompose plant residues can improve soil structure and enhance soil nutrient availability (Wu et al., 2012). In China, flue gas desulfurization gypsum by-product (FGDB) has been recognized as a new ameliorant for saline and sodic soils (Wang et al., 2012; Li et al., 2012). The responses of soil microbial communities to FGDB were examined using PLFA and real time quantitative polymerase chain reaction. Microbial community composition differed along the FGDB gradient; however, the microbial parameters did not follow a linear response. They found total PLFA concentrations, and concentrations of total bacterial and gram-negative bacteria PLFAs were slightly higher at intermediate FGDB concentrations. Similar results were obtained for the 16S rRNA copy numbers of bacteria at 0-20 cm depth. Shifts in PLFA-based microbial community composition and biomass could partly be explained by changes in soil physicochemical properties. Results conclude that addition of FGDB leads to significant changes in soil physicochemical and microbial parameters (Li *et al.*, 2012). Molecular phylogeny revealed the dominance of *Actinobacteria*, *Firmicutes* and *Proteobacteria* along with archaeal members of *Halobacteraceae* in saline-sodic soils (Keshri *et al.*, 2013). The library consisted of novel bacterial (20%) and archaeal (38%) genera showing \leq 95% similarity to previously retrieved sequences. Phylogenetic analysis indicated the ability of inhabitant to survive in saline and sodic conditions.

Microbial Enzymes in Saline and Sodic Soil

Soil enzymes are, generally, considered very sensitive to changes in soils and have been proposed as indicators to measure degree of soil stresses (Schimdt *et al.*, 2011; Nannipieri *et al.*, 2012). In spite of the extensive bibliography on the effects of soil use and management on potential soil enzyme activity, a very few studies have been focused on the effects of salinization and sodication on the enzyme activities. It is therefore of great interest to investigate how soil enzymes are affected by salinity and sodicity and which processes are responsible for their altered activities. Soil salinity and sodicity affect the activity of soil enzymes through its direct influence on microbial enzymatic production and structural changes in enzymes due to anionic movement and availability of organic matter (Frankenberger & Bingham, 1982; Amato & Ladd, 1992; Yao *et al.*, 2009).

An increase in soil salinity or sodicity inhibits the activity of dehydrogenase, β-glucosidase, urease, protease, alkaline phosphatase, acidic phosphatase, arylsulphatase, and argininamide hydrolyzing activity (Frankenberger & Bingham, 1982; Chander et al., 1994; Garcia & Hernandez, 1996; Rietz & Havnes, 2003; Tejada et al., 2006; Tripathi et al., 2007). Pathak & Rao (1998) reported that sodicity does not affect glutaminase and deaminase activities, while the activities of asparaginase and deaminase were not resistant to salinity and sodicity, respectively. This might be attributed to enzyme specific activity with respect to salinity and sodicity. Rietz & Haynes (2003) also reported the activities of β-glucosidase, alkaline phosphatase, and arylsulphatase decreased with increasing salinity (EC) and sodicity (ESP). The activity of arylsulphatase was less affected by salinity (EC) and sodicity (ESP). Inhibited activity of soil enzymes in saline soils appears to be associated with the physicochemical and microbial properties of soils (Frankenberger & Bingham, 1982). The decline in enzyme activity with increasing salinity also appeared to be related to change in osmotic potential of the soil because of higher salt concentrations and specific ion toxicity (Zahir et al., 2001). The NaCl (salt stress) induced changes in enzyme activities; this negative effect was even greater on hydrolases (phosphatase and β -glucosidase).

The activities of alkaline phosphatase and acidic phosphatase were higher in sodic (alkaline in old literature) and acidic soils, respectively (Tripathi *et al.*, 2007; Dick *et al.*, 2000; Singh *et al.*, 2012a, 2012b). Generally, alkaline phosphatase predominated in soils with neutral or slightly alkaline pH. It has been observed that activity of soil enzymes depends on microbial growth and availability of resources and nutrients (Sinsabaugh & Moorhead, 1994). The activity of alkaline phosphatase might be dependent on availability of labile phosphorus in the soils (Singh *et al.*, 2012a). The amount of inorganic P increased with increasing alkalinity and soil pH in sodic soils of semi-arid region. When soil microorganisms are phosphorus (P) limited, they produce acid or alkaline phosphatase (depending upon soil pH and microbial community composition) that release inorganic phosphate from organic matter (Allison *et al.*, 2011). Moreover, phosphatase activity has been shown to be inversely related to inorganic P availability in the soil system, especially in salt-affected soils (Allison *et al.*, 2007; Allison *et al.*, 2011). This relationship also holds at the global scale, where the ratio of P and carbon acquiring extracellular enzymes increases in tropical ecosystems; where P is more likely to limit productivity because of increased P weathering rates (Sinsabaugh *et al.*, 2008). Rao *et al.* (2000) reported that acid phosphatase–montmorillonite and acid phosphatase– tannate complexes had different pH-activity dependence and a higher and lower sensitivity to temperature and proteolysis, respectively. They concluded that phosphatase immobilized on clays, organic, and organo-mineral complexes showed catalytic features quite different from those of the free enzyme.



Figure 3. Relationship of dehydrogenase (DHA), alkaline phosphatase (Alkpa) and protease with soil pH [A, B, C] and exchangeable sodium percentage (ESP) [D, E, F] in a range of sodic soils. (*) P, 0.05; (**) P, 0.01. ESP, exchangeable sodium percentage. This figure is available in colour online at wileyonlinelibrary. com/journal/ldr

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Garcia et al. (1994) reported that reduction in enzyme activities may be attributed to lower microbial growth and biomass; thereby, less release in soil enzymes. In semi-arid soils, the enzyme activity is mainly extracellular, stable, and form complexes with the organic and mineral colloids (Schimdt et al., 2011; Nannipieri et al., 2012). Salt-tolerant microbes in salt-affected soils produce enzymes, which have optimum activities under higher salt concentration than corresponding enzymes produced by microbes in non-saline soils (Zahran, 1997). There are inverse relationships of soil pH and ESP with enzyme activities that generally show non-linear model (logarithmic, exponential, or polynomial) in a range of sodic soils. It was observed that protease enzyme decreased significantly with increasing soil pH, except alkaline phosphatase. While the dehydrogenase and protease showed a statistically significant negative relationship with ESP, alkaline phosphatase had significant positive relationship with ESP. The relationship of dehydrogenase with soil pH was weak (Figure 3). Soil pH has direct biochemical effects on activity of extracellular soil enzymes (Sinsabaugh et al., 2008). Soil pH reflects climatic controls on soil weathering and affects the large-scale distribution of soil enzymes through changes in nutrient availability (Sinsabaugh et al., 2008). The increase in EC disperses the clays, and the stable enzymes remain unprotected and therefore become more prone to denaturation (Frankenberger & Bingham, 1982; Pathak & Rao, 1998).

Garcia & Hernandez (1996) reported that NaCl was more effective than Na₂SO₄ in negatively affecting biological and biochemical fertility of soils, attributed to the toxic effect of particular ion on microbial growth in saline soils. Probably, chloride toxicity is greater than sulfate (Frankenberger & Bingham, 1982; Zahran, 1997; Pathak & Rao, 1998; Rietz & Haynes, 2003; Tejada et al., 2006). The addition of organic wastes had a positive effect on the activities of soil enzymes (Tejada et al., 2006). Lower values of enzyme activities in saline soils than in non-saline soils might be due to following two reasons: (i) 'salting-out' effect, which involves a decrease in enzyme solubility through dehydration, thus altering the enzyme 'catalytic site'; and (ii) enzymes in saline soils are less protected and perhaps they were denatured by proteolysis (Garcia et al., 2000). Recently, Pan et al. (2013) investigated that not all enzymes are sensitive to salinity. Urease was proved to have serious limitation with respect to its ability to reflect grassland salinization. Both β -glucosidase and alkaline phosphatase may be used as indicators of soil quality during grasslands salinization. Kang et al. (2013) assessed the changes and spatial distributions of soil enzyme activities in saline-sodic soil for 3 years of cultivation of Leymus chinensis under drip irrigation. Alkaline phosphatase, urease, and sucrase activities increased with cultivated years in saline-sodic soil under drip irrigation after planting L. chinensis for 3 years.

The available contradictory findings give generous justification for further investigations into the effects of salinity on C–N dynamics and structure, function, and diversity of soil microbial communities. I suggest following key research

issues to be addressed in future. What is the pattern of C-N dynamics in saline and sodic soils in relation to seasonal variations? How do salinity and sodicity affect N cycling, and will N dynamics follow the pattern similar to that of C dynamics? How is changing sodicity (in field conditions) likely to impact enzyme-mineral interactions, and will there be its any effect on C-N dynamics? Can we link diversity of mycorrhizal fungi and bacteria with aggregate stability in sodic soils? How we can manipulate these microbes to minimize C loss (CO2 emission) and maximize C stocks in saline and sodic soils? Are there distinct microbial (bacteria and fungus) species in relation to C-N dynamics at various salinity and sodicity levels in field soils? How dissolution of CaCO₃ (available in plenty in sodic soils) affects decomposition of organic matter, C-N mineralization, and microbial activities in sodic soils? How drying and rewetting of sodic soils from field, not in artificially created saline and sodic soils, affect C-N dynamics and microbial activities? What is the pattern of change in structure, function, and diversity of microbial communities during salinization and sodication? These issues are mysterious and need further investigations.

CONCLUSIONS

It is concluded that excess amounts of salts in soil not only have adverse effects on soil physical and chemical properties but also affect the microbial and enzyme activities. Both soil salinity and sodicity have detrimental effects on microbial communities (structure, function, and diversity). Drying and rewetting affect microbial activities of saline soils. The magnitude of its effect depends on salt concentration and availability of organic substrates in the soil. Because of structural instability, sodic soils are more prone to loss of soil organic carbon through C mineralization (CO2 emission). An increase in sodicity level tends to increase dispersion of clay particles, and this makes organic matter more accessible for microbial action. The microbial biomass, population, and values of potential soil enzyme activities are generally decreased with increasing salinity and sodicity of soil, which restrict the efficient N cycling.

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