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Paleocene-Eocene transition at Naqb Assiut, Kharga Oasis, Western Desert, Egypt: Stratigraphical and paleoenvironmental inferences

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#### Abstract

This work depends on the study of the lower part of the Esna Formation which 19 encompasses the Paleocene-Eocene (P-E) transition in Egypt as well as at Naqb 20 Assiut section, Kharga Oasis, Western Desert. The Paleocene/Eocene (P/E) 21 boundary is represented by El Dababiya Quarry Member which consists of five 22 distinctive beds (nos. 1-5) at the GSSP. On the other hand, at Naqb Assiut section 23 this boundary is only represented by the upper two beds (nos. 4&5), whereas, the 24 lower three beds (nos. 1-3) are missing due to a hiatus. This hiatus is marked by 25 the occurrence of an irregular surface contains pebbles and phosphatic materials. 26 This hiatus may be related to the echo of Sryian Arc Orogeny at the P/E time. 27 Biostratigraphically; four planktonic foraminiferal zones are defined from base to 28 top as: Acarinina soldadoensis/Globanomalina pseudomenardii and Morozovella 29 velascoensis (late Paleocene), Acarinina sibaiyaensis and Pseudohastigerina 30 wilcoxensis/Morozovella velascoensis (early Eocene). The Acarinina sibaiyaensis 31 Zone which represents the P//E/ boundary is characterized by the occurrence of 32 intrazonal hiatus at it's lower part. The benthonic foraminiferal taxa contain 33 abundant representatives of Midway-type fauna (~91 % of the whole 34 assemblages), beside few Velasco-type faunal ones (~9 %), indicating an outer 35 neritic (150-200 m) water depth of deposition during the P-E transition. 36 Quantitative analysis and composition of benthonic foraminiferal assemblages are 37 indicative for various environmental changes around the P/E boundary. They 38 reflected a high diversity, increase of epifaunal taxa, and low-intermediate 39 productivity conditions, which indicates a well-ventilated bottom water and oligo 40 - to mesotrophic conditions during the late Paleocene age. Rapid extinction of 41 about 18 % of the entire benthonic foraminiferal species started at the P/E 42 boundary, where the last occurrence of Angulogavelinella avnimelechi is 43 pronounced at the base of this boundary. There is a decline in the preceding faunal 44 characteristics of the late Paleocene and signified an increase in the abundance of 45 agglutinated and buliminid species. These characters are revealed some 46 environmental stressful conditions of low oxygen and increased productivity. 47 During the early Eocene, the environmental conditions are improved leading to 48 49 the recovery of most benthonic and planktonic foraminifera.

**Keywords:** Benthonic and planktonic foraminifera; P/E boundary; Midway-type fauna; 50 Velasco-type fauna; Paleobathymetry; biostratigraphy; paleoenvironment. 51

## **1. Introduction**

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During the last few decades, much researches have been focused on the 53 Paleocene-Eocene transition, on the study of the abrupt and temporal period of 54 extreme global warming event (Zachos et al., 1993) and its effects on the various 55 marine and terrestrial biota (Aubry et al., 1998). The planktonic foraminiferal 56 fauna show good diversification and faunal turnover after and before the P/E 57 boundary event (Canudo & Molina, 1992; Lu & Keller, 1995; Canudo et al., 58 1995; Arenillas & Molina, 1996; Lu et al., 1996; Arenillas et al., 1996). Also, the 59 P/E boundary was marked by the mass extinction of benthonic foraminifera 60 (Thomas, 2007), which is commonly known as the Benthonic Extinction Event 61 (Tjalsma, 1977; Schnitker, 1979; Tjalsma and Lohmann, 1983; Thomas, 1990, 62 2003). The Global Stratotype Section and Point (GSSP) of the P/E boundary was 63 defined at El Dababiya Village, south Luxor, Egypt. The P/E boundary at the 64 GSSP was located at the base of five characteristic beds known as Dababaiya 65 Quarry beds within the Esna Formation (Dupuis et al., 2003). Aubry et al. (2007) 66 raised the stratigraphic rank of these beds into a member (e.g. El Dababiya Quarry 67 Member). The base of this distinctive member is equivalent to the Paleocene 68 Eocene Thermal Maximum (PETM) and the base of Carbon Isotope Excursion 69 (CIE). It is characterized by the extinction of Angulogavelinella avnimelechi, the 70 influx of planktonic foraminiferal excursion taxa (e.g. Acarinina sibaiyaensis, A. 71 Africana, Morozovella allisonensis) and that of the Discoaster araneus-72 Rhomboaster spp. assemblage (Dupuis et al., 2003). 73

Kharga Oasis is one of the most common Egyptian depressions in the Western 74 Desert, which lying between Lat. 24° and 26° N, with about 200 km to the west of 75 the Nile Valley (Fig. 1). It is bounded by a north - south western side of an 76 irregular and continuous escarpment, where the Naqb Assiut section is located at 77 the northern part of this scarp, and opened from the southern and south western 78 parts. At Kharga Oasis, several studied were carried out on the Paleocene-Eocene 79 rocks such as Hewaidy, 1983; Luger, 1985; Tantawy, 1998; Mahfouz, 2008; El- 80 Azabi and Farouk, 2010; Khalil and Al Sawy, 2014. These studies were 81

concentrated on the general stratigraphy of the Kharga Oasis, except for Mahfouz 82 (2008) who was focused his study on the P/E boundary. 83

To investigate the P/E lithological boundary, study the planktonic and benthonic 84 foraminiferal fauna and their extinction across the P-E transition, high resolution 85 study is carried out on the Naqb Assiut section. Qualitative and Quantitative 86 analyses of these planktonic and benthonic fauna will be documented here to 87 show; the planktonic foraminiferal zones, the relative abundances of their species 88 across the extinction level. Moreover, the uses of the benthonic foraminiferal 89 morphogroups (Corliss, 1985) are used to deduce paleoxygenation conditions of 90 bottom water environment. Also, the foraminiferal indices (e. g.; genus and 91 species richness, Fisher  $\alpha$  – diversity,.....etc.) are used to detect the nature of the 92 P/E boundary and infer the paleoenvironmental changes throughout the P-E 93 transition of the present area.

### 2. Materials and methods

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A total of eighteen samples have been collected from the Esna Formation at Naqb 96 Assiut section with close sample spacing around the boundary. About 200 g of dry 97 rock/sediment were disaggregated in water with diluted hydrogen peroxide 98 ( $H_2O_2$ ). Afterwards, samples were washed through a sieve with 63 µm, and the 99 obtained residues were dried. This procedure was repeated until foraminifera with 100 clean surface texture were recovered. Population counts follow the method of 101 Buzas (1990) and were based on random splits of at least 300 specimens 102 (Table1).

Quantitative studies of planktonic and benthonic foraminiferal species are 104 depended mainly on representative splits of about 300 specimens >  $63\mu$ m (Table 105 1). All representative specimens are picked and mounted on microslides for 106 identification and a permanent record. Studying the benthonic foraminiferal 107 morphogroup analysis (e. g. Corliss, 1985; Jones and Charnock, 1985; Corliss and 108 Chen, 1988) allows deducing probable microhabitat preferences and 109 environmental conditions (Table 2) related to sea-water oxygenation (e. g., 110 Bernhard, 1986; Jorissen *et al.*, 1995; Kaminski and Gradstein, 2005). It 111

expressed here as the epifaunal and infaunal morphotypes (E/I %). Relative 112 abundances of most common benthonic taxa (Fig. 8), calcareous to agglutinated 113 ratios (C/A %), and other proxies referred to diversity such as; the Fisher  $\alpha$  index 114 (Murray, 1991), in addition to both of the distinct depth-controlled benthonic 115 fauna of the Midway-type and the Velasco-type faunal ratios (MF/VF %,) 116 (Berggren and Aubert, 1975).

The comparison between fossils and recent benthonic foraminiferal data, the 118 occurrence and abundance of depth-related species and the use of their upper 119 limits (Van Morkhoven *et al.*, 1986; Alegret *et al.*, 2001, 2003) allowed inferring 120 the paleodepth of these assemblages. Moreover, ratios of most important 121 planktonic foraminiferal groups such as the acarininids, morozovellids and 122 subbotinids are calculated (Fig. 9). Microslides, rock-samples and residues are 123 deposited in the Geological Museum of the Geol. Depart., Fac. Sci., Minia 124 University.

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#### 3. Lithostratigraphy:

The P-E transition generally lies within the lower part of Esna Formation in 127 Egypt. The Esna Formation was first described by Beadnell (1905) to define the 128 shale succession between Duwi Formation at the base and Thebes Formation at 129 the top. Said (1962) restricted the Esna Formation of Beadnell (1905) into the 130 shale succession from Tarawan Formation (chalky limestone) at the base to the 131 Thebes Formation (Limestone) at the top. Abdel-Razik (1972) classified the Esna 132 Formation into two members, El Hanadi Member at the base and El Shaghab 133 Member at top. Also, He (op,cit) proposed Abu Had Member for the lower part of 134 the Thebes Formation. Recently, Aubry et al., (2007) modified the classification 135 of Abdel-Razik (1972) and subdivided the Esna Formation into four members 136 from base to top: El-Hanadi, El Dababiya Quarry, El-Mahmiya and Abu Had (Fig. 137 4) and they considered the last member as a member of Esna Formation rather 138 than Thebes Formation. At the study section, the lower part of the Esna Formation 139 was sampled and described (total thickness about 10.70 m, Figs. 2 & 3). It 140 represents El-Hanadi, El Dababiya Quarry and the lower part of El-Mahmiya 141 members. These members will be discussed briefly as follows: 142

#### 3.1. El-Hanadi Member (7.70m):

This member was originally introduced by Abdel Razik (1972) to describe the 144 lower part of Esna Formation (5 meters of shale succession) overlies the Tarawan 145 Formation and underlies El Shaghab Member at Gabal El Shaghab, east of El 146 Hanadi Village near Esna City. El Hanadi Member has been emended by Aubry *et* 147 *al.* (2007) to be coeval the lower part of the Esna Formation below El Dababiya 148 Quarry Member. It is equivalents to the Esna Unit 1 of Dupuis *et al.*, (2003) (Fig 149 4).

At Naqb Assiut section, El-Hanadi Member consists of about 7.70 m. thick. 151 (samples 1- 12), grayish green, papery shale (~ 6.75 m., samples 1-9) at base, 152 which varies gradually into massive marl (0.95 m. thick., samples no. 10-12) at 153 top (Figs. 2 & 3).

## **3.2. El Dababiya Quarry Member** (0.90 m.): 155

This Member was originally introduced by Aubry *et al.* (2007) as El Dababiya 156 Quarry Member instead of El Dababiya Quarry Beds of Dupuis *et al.*, (2003) in 157 the Global Standard Stratotype section and Point (GSSP) of the P/E boundary at 158 the Village of El Dababiya. El Dababiya Quarry Member constitutes a distinctive 159 lithologic succession of five characteristic beds (1-5) of about 3.68 m thick at it's 160 type locality. Bed no. 1 is a dark clay (~0.63 m thick); bed no. 2 is a phosphatic 161 brown shale (~ 0.50 m thick); bed no. 3 is a creamy phosphatic shale (~ 0.84 m 162 thick); bed no. 4 is a grey calcareous shale (~ 07.1 m thick) and bed no. 5 is a 163 marly calcarenitic limestone (~ 1.00 m thick). At Naqb Assiut section, El 164 Dababiya Quarry Member is only represented by the upper two beds (nos. 4&5) 165 (Figs. 2-5) from base upward as: 166

- Bed no. 4: It attains about 20 cm thick. It is massive marl with phosphatic 167 materials and pebbles at base. The lower surface of this bed is irregular and 168 bioturbated. It is coeval to bed no. 4 of the same member at the GSSP section and 169 is here representing by sample no. 13. The criteria at the base of this bed indicate 170 the occurrence of a hiatus at the P/E boundary.

Bed no. 5: It is about 70 cm. It is a yellowish chalky limestone and is coeval to 172 bed no. 5 of El Dababiya Quarry Member at the GSSP section. This bed is here 173 representing by samples no. 14 and 15.
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#### 3.3. El-Mahmiya Member (2.00m.):

This member was originally defined by Aubry *et al.* (2007) to cover the Esna Unit 176 2 of Dupuis *et al.*, (2003) at El Dababiya Village. The lower part of this member 177 (~ 2m thick) is only considered in this study. It consists of marly shale, and it is 178 represented by samples no. 16 -18.

#### 4. Planktonic Foraminiferal Biostratigraphy:

According to the important planktonic foraminiferal species, four zones for the P- 181 E transition are recognized at Naqb Assiut section (Figs. 6&7). The planktonic 182 foraminiferal zonal scheme of Berggren and Pearson, (2005) and it's categories of 183 zones are here applied. The proposal planktonic foraminiferal zones arranged 184 from base to top as follows: 185

#### **4.1.** *Acarinina soldadoensis/Globanomalina pseudomenardii* Zone 186

This zone was defined as Concurrent-Range Zone by Berggren *et al.* (1995) to 187 cover the interval from the Lowest Occurrence (LO) of *A. soldadoensis* 188 (Brönnimann) to the Highest Occurrence (HO) of the *G. pseudomenardii* (Bolli). 189 It represents the greater part of El-Hanadi Member (~ 6.75 m. thick, samples 1-9). 190 The *A. soldadoensis/G. pseudomenardii* (P4c) Zone is conformably overlain by 191 the *Morozovella velascoensis* (P5) Zone at the study section. 192

## 4.2. Morozovella velascoensis Zone

*M. velascoensis* Zone was defined as Partial Range Zone by Berggren & Pearson 194 (2005) to cover the interval from the HO of *G. pseudomenardii* (Bolli) to the LO 195 of the *Acarinina sibaiyaensis* (El Naggar). Berggren & Ouda (2003 a, b) 196 subdivided P5 Zone (sense of Berggren *et al.*, 1995) into three subzones namely: 197 P5a (latest Paleocene), P5b to recognize the Carbon Isotope Excursion (CIE) and 198 Paleocene/Eocene Thermal Maximum (PETM) interval and P5c to recognize the 199 post CIE and PETM interval (earliest Eocene). Berggren & Pearson (2005) used 200 P5 Zone (=P5a Subzone of Berggren & Ouda, 2003 a,b) to recognize the latest 201 Paleocene and used E1 (=P5b Subzone of Berggren & Ouda, 2003 a,b) to 202

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recognize the CIE & PETM interval and E2 Zone (=P5c Subzone of Berggren & 203 Ouda, 2003 a,b) to recognize the post CIE & PETM interval (Fig. 7). 204

In the present work, the *M. velascoensis* (P5) Zone is defined as a partial range 205 zone from the HO of *G. pseudomenardii* (Bolli) to the LO of Planktonic 206 Foraminiferal Excursion Taxa such as: *A. africana* (El Naggar) and *A.* 207 *sibaiyaensis* (El Naggar) at the P/E boundary. It represents the uppermost part of 208 El-Hanadi Member (~ 1 m. thick, samples 10-12). *M. velascoensis* (P5) Zone is 209 unconformably overlain by the *A. sibaiyaensis* (E1) Zone due to the missing of the 210 lower part bed 1-3 of El Dababiya Quarry Member. This zone is equivalent to P5 211 Zone of Berggren & Pearson, (2005) (Fig. 7).

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#### 4.3. Acarinina sibaiyaensis Zone

This zone was originally defined by Molina *et al.* (1999) to cover the interval 214 from the LO of the nominate taxon to the LO of *Pseudohastigerina wilcoxensis* 215 (Cushman & Ponton). It is here defined as a Lowest Occurrence zone according to 216 the definition of Molina *et al.* (1999). The *A. sibaiyaensis* Zone represents the 217 main part of El Dababiya Quarry Member (~0.90 m. thick, samples 13-15). The 218 lower part of this zone is missing due to the absence of the lower three beds of El 219 Dababiya Quarry Member. *A. sibaiyaensis* (E1) Zone is conformably overlain by 220 the *Pseudohastigerina wilcoxensis/Morozovella velascoensis* Zone. This zone is 221 equivalent to the upper part of El Zone of Berggren & Pearson, (2005) (Fig. 7).

#### **4.4.** *Pseudohastigerina wilcoxensis/Morozovella velascoensis* Zone 223

It was originally defined as Concurrent-Range Zone by Berggren and Ouda (2003) 224 from the LO of *P. wilcoxensis* (Cushman & Ponton) to the HO of *M. velascoensis* 225 (Cushman). It is completely equivalent to *P. wilcoxensis* Zone of Molina *et al.* 226 (1999), the upper part of *M. velascoensis* (P5) Zone of Berggren *et al.* (1995) (Fig. 227 7) and E2 Zone of Berggren & Pearson, (2005). It represents the lower part of El- 228 Mahmiya Member (~ 2.0 m. thick, samples 16-18). 229

## 5. Paleobathymetry

Benthonic foraminiferal assemblages of the upper Paleocene sediments at Naqb 231 Assiut section are cosmopolitan and dominated by elements of the so-called 232 Midway-type fauna (MF,  $\sim 93\%$  of the whole assemblages), typical of outer 233 neritic water depths (150  $\sim 200$  m), beside few deeper-water (upper bathyal) 234

Velasco-type faunal elements (VF, ~ 7 %). The Midway-type faunal elements 235 include outer neritic species such as; Bulimina quadrata, Cibicidoides alleni, C. 236 succedens, Gyroidinoides girardana, G. subangulata, Lenticulina midwayensis, L. 237 pseudomamilligera, Loxostomoides applinae, Marginulinopsis tuberculata, 238 midwayensis, Siphogenerinoides eleganta, *Stilostomella* Valvulineria 239 scrobiculata, Bulimina farafraensis, Oridorsalis plummerae, Spiroplectinella 240 esnaensis together with some nodosariids and others (Speijer, 1994; Alegret et al., 241 2005). Upper bathyal taxa (VF) are encountered with low numbers within El 242 Hanadi Member and comprises of Angulogavelinella avnimelechi, Anomalinoides 243 rubiginosus, Neoflabellina jarvisi, Gaudryina pyramidata and others. It is 244 interesting to find the A. avnimelechi in the lower part of the section (with few 245 specimens), which is common in the Paleocene outer neritic to upper bathyal 246 deposits in Tethyan and European basins (Speijer et al., 1995; Speijer and 247 Schmitz, 1998; Dupuis et al., 2003; Alegret et al., 2005; Ernst et al., 2006). Also, 248 the lower Eocene sediments are characterized by the scarcity of VF ( $\sim 5\%$ ) and 249 the abundance of MF ( $\sim 95\%$ ) as well as the dominance of opportunistic faunas 250 (e. g., Valvulineria scrobiculata, Anomalinoides egyptiacus, A. zitteli, Stainforthia 251 farafraensis and Bulimina farafraensis) at the P/E boundary, which are thrived in 252 the shallow-outer neritic environment (Speijer, 1994; Speijer and Schmitz, 1998). 253 These indicate that, the Naqb Assiut area was not deeper than 200 m. during the 254 early Eocene age. Consequently, the stratigraphic distribution of most of the 255 Paleocene – Eocene benthonic assemblages and the presence of both elements 256 (MF and VF) suggest an outer neritic (~ 150-200 m. water-depth) environment. 257

## 6. Foraminiferal associations

The Paleocene benthonic foraminiferal assemblages of the Esna Formation at 259 Naqb Assiut section (lower part 7.70 m. thick., El Hanadi Member, samples 1-12) 260 are diverse and heterogeneous ( $\alpha$ = 26.8, H (S)= 4.1, respectively, Fig. 9). 261 Calcareous foraminifera represent the major component of the assemblage (75 – 262 85 %), while the agglutinated forms are minor ones (15 – 25 %). Among the 263 calcareous taxa; *Loxostomoides applinae, Cibicidoides alleni, Osangularia* 264 *plummerae, Bulimina midwayensis, Lenticulina midwayensis,* and *L. navicula* are 265 most abundant (Fig. 9). Benthonic assemblages consist of a mixture of epifaunal 266

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(42 - 60 %) and infaunal (40 - 58 %) morphogroups. Planktonic foraminiferal 267 assemblages within the Acarinina soldadoensis/Globanomalina pseudomenardii 268 (P4c) and Morozovella velascoensis (P5) zones are diverse and rich in tropical - 269 subtropical species. The Neoflabellina jarvisi, Angulogavelinella avnimelechi, 270 pseudomamilligera, Laevidentalna Lenticulina gracilis and Discorbis 271 pseudoscopos (Fig. 9 and Table 1) disappeared at the P/E boundary. The 272 agglutinated individuls of Spiroplectinella dentata and pseudoclavulinids are 273 increased somewhat in abundance. On the other hand, the Clavulinoides asper 274 whitei and Gaudryina pyramidata are disappeared at the P/E boundary. The 275 diversity and species richness ratios of the assemblages reached immediately its 276 acme below the P/E boundary (sample no. 12), where the epifaunal morphogroups 277 decreased. They are exactly declined at the P/E boundary. The infaunal 278 morphogroups (Table 2) are increased at the P/E boundary (~79 % at the base of 279 El Dababiya Quarry Member, sample no. 13). The lowermost part of El Dababiya 280 Quarry Member (sample no. 13, Figs. 8 & 9) is almost barren of benthonic 281 foraminifera, whereas planktonic foraminifera are common. They are dominated 282 by some agglutinated taxa of mainly pseudoguadryinids, spiroplectamminids and 283 verneuilinids, in addition to few specimens of nodosariids (Table 1). The 284 acarininids dominate the planktonic foraminiferal assemblages (~ 43%), relative 285 to the morozovellids and subbotinids (~ 11% and ~ 28%, respectively). 286

In the upper part of the El Dababiya Quarry Member, the benthonic foraminiferal 287 assemblages are recovered, where the percentages of diversity and species 288 richness of the assemblages (Fig. 9) are gradually increased towards the 289 uppermost part of the member. Assemblages are dominated by the benthonic 290 calcareous foraminifera (~83%) of mainly bagginid and stainforthiid species, such 291 as; *Valvulineria scrobiculata, Stainforthia farafraensis, Globocassidulina* 292 *subglobosa, Anomalinoides zitteli* and *A. aegyptiacus*. Here, the lenticulinids (~ 293 43 %) are replaced by the buliminids (~57 %), and include *Bulimina farafraensis,* 294 *B. midwayensis and B. esnaensis* (Figs. 8 & 9 and Table 1).

At the uppermost part of the section (El Mahmiya Member, samples no. 16-18, 296 Fig. 9), in spite of the increase of calcareous taxa (~90%), the diversity and 297 species richness are decreased (Fig. 9). Acarininids maintained with its relative 298

abundance (~ 37 %), in regard to the morozovellids and subbotinids (~ 22% and ~ 299 30%, respectively, Fig. 9). Also, the buliminids are still dominated in this zone 300 with ~ 65 % against ~ 35% for the lenticulinids (Fig. 9). 301

### 7. Paleoenvironmental inferences

The quantitative analyses of the benthonic foraminiferal assemblages allow us to 303 explain the foraminiferal distribution pattern of studied section across the P/E 304 boundary. 305

The upper Paleocene benthonic foraminiferal assemblages in Nagb Assiut (El 306 Hanadi Member, samples 1-12, Fig. 9) are diverse and composed mainly of 307 Midway-type faunas (MF average of  $\sim 93$  % of total assemblages), against a 308 minor Velasco-type ones (VF  $\sim$  7%, Fig. 9). It has a high foraminiferal number 309 (with average of ~ 1703 specimens in the 125  $\mu$ m size fraction per gram dry 310 sediments), Fisher  $\alpha$ - diversity index (average of  $\sim 26.8$ ) and Species richness (65 311 species, Fig. 9). This interval contains a mixed morphogroups (Corliss, 1985, 312 1991; Corliss and Chen, 1988; Widmark, 1997; Culver, 2003) of more epifaunal 313 taxa (~ 52%) than infaunal ones (~ 49%), suggesting a well ventilated, normal 314 marine and oligo - to mesotrophic levels. Presence of mixed epi-and infaual 315 morphotypes suggests less abundant food resources to benthos (Jorissen et al., 316 1995; Van der Zwaan et al., 1999) or living in mesotrophic conditions (Gooday, 317 2003). This interval is dominated by Loxostomoides applinae, Spiroplectinella 318 dentata, Cibicidoides alleni and Lenticulina navicula. The presence of calcareous 319 thick-walled taxa, such as Cibicidoides alleni, C. pseudoacutus and Lenticulina 320 midwayensis indicates high oxygen levels. According to Berhnard (1986) and 321 Speijer and Wagner (2002), the Gyroidinoides and Lenticulina are epifaunal and 322 oxic bottom water genera (plano-convex or lenticular morpho groups). Moreover, 323 the upper Paleocene benthonic foraminiferal assemblages are enriched and 324 dominated by the Lenticulina spp. (~ 66% of the assemblages), on the contrary of 325 the buliminids ( $\sim 34\%$ ). This confirmed the well-oxygenated bottom waters or 326 low to intermediate productivity conditions. With respect to the planktonic 327 foraminifera, Acarinina, Morozovella and Subbotina make up 88 % of the total 328 planktonic assemblage in this interval (Fig. 9). Frequencies of Igorina, 329

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*Parasubboina* and *Globanomalina* form a relative percentage less than 12%. So, 330 the benthonic foraminiferal assemblages of the upper Paleocene sediments suggest 331 an outer neritic environment (~ 150-200 m water depth), associate with oligo- to 332 mesotrophic levels and aerobic conditions at the sea floor with low to intermediate 333 productivity. 334

The P/E boundary is recorded at the base of a  $\sim 20$  cm thick phosphatic marl bed. 335 The lower surface of this bed is irregular, bioturbated, and contians pebbles and 336 phosphatic materials, which indicates a short period of hiatus at the P/E boundary 337 (Fig. 9). This bed contains remarkable warm taxa, suchas; Acarinina sibaiyaensis 338 and A. africana. At the same time, the agglutinated taxa are represented very well. 339 The benthonic foraminiferal distribution pattern (Table 1) revealed generally that 340 89% of MF species (e.g. Stilostomella midwayensis, Gavelinella danica, Tritaxia 341 applinae, *Coryphostoma* midwayensis, Loxostomoides midwayensis, 342 Spiroplectinella dentata, Cibicidoides alleni, Siphogenerinoides eleganta, 343 Gyroidinoides subangulata, Tappanina selmensis, Bulimina midwayensis, 344 Gaudryina textulariformis, Anomalinoides aegyptiacus, Pyramidulina vertebralis, 345 Eponides lunatus), and 6 % of VF ones (e.g. Clavulinoides amorpha, C. trilatera, 346 Anomalinoides rubiginosus, Bulimina trinitatensis, Marssonella oxycona 347 oxycona, M. oxycona trinitatensis) (Berggren and Aubert, 1975) occur above the 348 P/E boundary. On the other hand, 5% of VF species extinct near the same 349 boundary (e.g. Bolivinopsis spectabilis, Neoflabellina jarvisi, Clavulinoides asper 350 whitei, Gaudryina pyramidata, Angulogavelinella avnimelechi) (Fig. 9, and Table 351 1). Within this bed, there is a relative increasing in the infaunal morphogroup 352 (~79%) as well as in the agglutinated foraminiferal taxa (Fig. 9). This interval has 353 a drop in the foraminiferal number (~1200 specimens in the 125 µm size fraction 354 per gram dry sediments). Accordingly, the deposition of this interval may be in 355 middle-shallow outer neritic setting (50-150m water depth) associated with 356 oceanic environmental stresses probably related generally to moderate trophic and 357 oxygen deficiency conditions. 358

The upper part of El Dababiya Quarry Member consists of chalky limestone bed 359 of 0.70 m thick (samples 14-15). It contains high relative percentage of both 360 planktonic and benthonic foraminiferal species. Within this bed, the 361

environmental conditions are improved with enough nutrients and oxygen content. 362 There are an increasing in the foraminiferal faunal number (~ 1362 specimens in 363 the 125  $\mu$ m size fraction per gram dry sediments), Fisher  $\alpha$ - diversity (Fig. 9). 364 Also, it is dominated by the MF taxa (~ 97 % MF against 3 % for the VF). Also, 365 the agglutinated taxa are marked by decreasing in their relative abundance. 366 Moreover, the buliminids (~ 57 %) are exceeded the lenticulinids (~ 43%) (Fig. 367 9). At the same time, the planktonic foraminiferal assemblages are almost 368 maintained its faunal characteristics, where their main elements of Morozovella, 369 Acarinina, and Subbotina are making up 84%, while the remainder ones (Igorina, 370 Parasubboina, Pseudohastigerina and Globanomalina) are reached ~ 16% of the 371 whole planktonic assemblages. This interval is dominated by the Valvulineria 372 scrobiculata, Anomalinoides aegyptiacus, A. zitteli, and Stainforthia farafraensis 373 of outer neritic setting (paleodepth 150-200m). Accordingly, the deposition of this 374 interval may reflect an outer neritic shelf environment, associated with high 375 productivity and increased food supply conditions. 376

The topmost part of the Naqb Assiut section is represented by a marly shale bed of 377 about 2.00 m. thick (samples 16-18) which is belonging to the El Mahmiya 378 Member. It witnessed somewhat low diversity values in both planktonic and 379 benthonic foraminiferal taxa. This interval is characterized by a foraminiferal 380 number fluctuates between 1196 and 1674 specimens in the 125  $\mu$ m size fraction 381 per gram dry sediments. Here, the MF taxa are continued almost with the same 382 abundance of the preceding interval (~97 % of the whole assemblage), while the 383 VF ones are represented with few numbers (Fig. 9). Buliminids and lenticulinids 384 taxa make up ~65% and ~35% of the relative abundance of buliminids 386 is thought to be due to the abundance of food supply (Fontanier *et al.*, 2002; 387 Gooday, 2003; Alegret and Thomas, 2004). Moreover, the infaunal morphogroups 388 are here increased and reached ~ 61% (Fig. 9), on the expense of the epifaunal 389 ones (~ 39%). This indicates a high nutrient flux to the sea floor.

Benthonic foraminiferal assemblages of this interval are characterized by the 391 abundance of *Stainforthia farafraensis*, *Valvulineria scrobiculata*, then 392 *Anomalinoides* spp. (e.g., *A. aegyptiacus, A. zitteli*, and *A. umbonifera*) and 393

*Bulimina farafraensis*. According to Speijer (1994) and Speijer and Schmitz 394 (1998), the presence of *V. scrobiculata, S. farafraensis, A. aegyptiacus, A. zitteli* 395 indicates an outer neritic setting (~150-200 m), associated with high food supply 396 and increased productivity conditions. 397

## 8. Conclusions

Similar to the GSSP of the P/E boundary at Dababiya Village, south of Luxor 399 (~35), Egypt, the P/E boundary at Naqb Assiut section lies at the base of El 400 Dababiya Quarry Member within the lower part of the Esna Formation. El 401 Dababiya Quarry Member consists of five distinctive beds (1-5) at the GSSP 402 locality. On the other hand, at the study section, the lower beds of El Dababiya 403 Quarry Member (1-3) are missing and the upper beds are only represented (4and 404 5).

The lower surface of the bed no. 4 is irregular, bioturbated and contains pebbles 406 and phosphatic materials. The bed (4) attains about 20 cm thick, but the bed (5) 407 attains about 70 cm thick. The correlation of the beds of El Dababiya Quarry 408 Member at the GSSP and Naqb Assiut section indicates the occurrence of a hiatus 409 at Naqb Assiut. This hiatus is related to a syn-sedimentary tectonic event at the 410 P/E boundary, which may be related to the echo of Syrian Arc Orogeny in Egypt 411 during this time.

Biostratigraphically, four planktonic foraminiferal zones are identified for the P/E 413 transition from base to top: *Acarinina soldadoensis/Globanomalina* 414 *pseudomenardii* (P4c) and *Morozovella velascoensis* (P5) for the late Paleocene, 415 *Acarinina sibaiyaensis* (E1) and *Pseudohastegerina wilcoxensis/Morozovella* 416 *velascoensis* (E2) for the early Eocene. The P/E boundary is defined at P5/E1 417 zonal boundary, with intarzonal hiatus at the base of E1 Zone (early Eocene). 418

The benthonic foraminiferal assemblages as well as their quantitative analyses 419 indicate generally a deep outer neritic setting (150-200 m water depth) of 420 deposition during the late Paleocene and early Eocene age. Characteristics of the 421 different P-E faunal assemblages are discriminated into the following: 422

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-The late Paleocene faunal assemblages of El Hanadi Member were highly 423 diverse. It is mainly dominated by the *Loxostomoides applinae*, *Cibicidoides* spp. 424 and *Lenticulina* spp. These taxa indicate high oxygen levels, oligo- to 425 mesotrophic, and low to intermediate productivity conditions. 426 Paleobathymetrically, most of these benthonic faunal taxa indicate outer neritic 427 environment (150-200 m).

-At the P/E boundary (base of El Dababiya Quarry Member, phosphatic marl 429 bed), benthonic foraminiferal assemblages exhibit a major faunal turnover due to 430 the occurrence a hiatus. This interval contains agglutinated benthonic 431 foraminiferal assemblages. The marl bed of El Dababiya Quarry Member may be 432 attributed to the occurrence of upwelling with nutrient-rich deep water. It is 433 marked by an oceanic dysoxic conditions, this is documented by the occurrence of 434 phosphatic materials increasing in the relative abundance of agglutinated 435 foraminiferal taxa. Also, it is marked by the lowest diversity in the infaunal taxa. 436 The benthonic foraminiferal fauna indicate a middle shallow-outer neritic setting 437 (50-100 m water depth).

-Within the upper part of El Dababiya Quarry Member (chalky limestone bed), 439 there is increasing in the diversity of both planktonic and benthonic foraminiferal 440 with respect to the lower part. It witnessed some environmental changes that led 441 to the recovery of most benthonic foraminiferal assemblages. The *Anomalinoides* 442 *aegyptiacus* and abundant buliminids are indicative to stress environmental 443 condition with low oxygen levels and high productivity. The benthonic 444 foraminiferal taxa indicate a shallow outer neritic setting (100 m water depth).

-During the early Eocene (El Mahmiya Member) there are a relative decreasing in 446 the diversity values of the benthonic foraminiferal fauna. *V. scrobiculata, S.* 447 *farafraensis* and buliminids dominated the benthonic foraminiferal assemblages, 448 which are thought to be representatives of high food and increased productivity 449 conditions. The benthonic foraminferal taxa indicate an outer neritic setting (150- 450 200 m water depth).

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## **Figure Caption**

Fig. 1. Geological map of Kharga Oasis shows the location of Naqb Assiut 659outcrop, Western Desert, Egypt (modified after Said, 1990).660

Fig. 2. A Field photograph shows the P/E boundary at Naqb Assiut section, B. 661 Field photograph shows pebbles, collophane and coprolite grains at the basal part 662 of El Dababiya Quarry Member, C. Field photograph shows bioturbation at base 663 of El Dababiya Quarry Member. 664

Fig. 3. Lithostratigraphic columnar section of the P-E succession at Naqb Assiut 665 section, Western Desert, Egypt. 666

Fig. 4. Lithostratigraphic correlation for the P-E transition rocks at the present 667 work with other pervious ones. 668

Fig. 5 Lithostratigraphic correlation between El Dababiya Quarry Member at the 669GSSP and Naqb Assiut sections.670

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Fig. 6. Biostratigraphic distribution chart of the planktonic foraminiferal species 671 recorded throughout the P-E succession at Naqb Assiut section, Western Desert, 672 Egypt. 673

Fig. 7. Comparison between the present planktonic foraminiferal zones with some 674 local and international ones. 675

Fig. 8. Stratigraphic distribution and relative abundance of most common 676 benthonic foraminiferal taxa and its turnover across the P/E boundary (dashed 677 line) at Naqb Assiut section. Taxa whose abundance is  $\geq 2$  % in at least one 678 sample have been plotted. For other taxa see Table (1). 679

Fig. 9. Benthonic foraminiferal indices, percentages of planktonic foraminiferal 680 *Acarinina* and *Morozovella*, and paleoenvironmental inferences across the 681 Paleocene/Eocene boundary at Naqb Assiut section, Western Desert, Egypt. 682

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# **Plate Captions**

Plate1. (Scale bar is 100 μm); 1-2 Globanomalina pseudomenardii (Bolli), 685sample no. 5; 3-4 Morozovella velascoensis (Cushman), sample no. 18; 5-6 686Morozovella occlusa (Loeblich & Tappan), sample no. 8; 7-8 Acarinina africana 687(El Naggar), sample no. 14; 9-10 Acarinina sibaiyaensis (El Naggar), sample no. 68814; 11-12 Acarinina soldadoensis (Brönnimann), sample no. 2; 13-14 Subbotina 689triloculinoides (Plummer), sample no. 6; 15-16 Pseudohastigerina wilcoxensis 690(Cushman & Ponton), sample no. 17.

**Plate 2.** (Scale bar is 100 µm); 1 *Ammodiscus glabratus* Cushman and Jarvis, 692 Sample no. 1; 2 *Ammosphaeroidina pseudopauciloculata* (Myatlyuk), sample no. 693 3; 3 *Bolivinopsis spectabilis* (Grzybowski), sample 2; 4 *Spiroplectinella dentata* 694 (Alth), sample no. 2; 5 *Ammoglobigerina altiformis* (Cushman and Renz), sample 695 no. 6; 6 *Karrerulina conversa* (Grzybowski), sample no. 13; 7 *Gaudryina* 696 *pyramidata* Cushman, sample no. 6, 8 *Bermudezina danica* (Franke), sample no. 697 1; 9 *Tritaxia midwayensis* (Cushman), sample no. 3; 10 *Dorothia bulletta* 698 (Carsey), sample no. 4; 11 *Marssonella oxycona oxycona* (Reuss), sample no. 2; 699

12 Calvulinoides amorpha (Cushman), sample no. 7; 13 Calvulinoides trilatera 700 (Cushman), sample no. 11;14 Pseudoclavulina farafraensis LeRoy, sample no. 2;701 15 Spiroloculina proboscidae Schwager, sample no. 3; 16 Laevidentalina colei 702 (Cushman and Dusenbury), sample no. 4; 17 Nodosaria longiscata d'Orbigny, 703 sample no. 3; 18 Pyramidulina latejugata (Gumbel), sample no. 4; 19704 Frondicularia goldfussi Reuss, sample no. 1; 20 Lenticulina midwayensis 705 (Plummer), sample no. 2; 21 Lenticulina navicula (d'Orbigny), sample no. 5; 22 706 Lenticulina pseudomamilligera (Plummer), sample no. 6; 23 Marginulinopsis 707 tuberculata (Plummer), sample no. 5; 24 Saracenaria triangularis (d'Orbigny), 708 sample no. 7; 25 Neoflabellina paleocenica Titova, sample no. 3; 26 Lagena 709 hispida Reuss, sample no. 4; 27 Glandulina laevigata (d'Orbigny), sample no. 6; 710 28 Loxostomoides applinae (Plummer), sample no. 2; 29 Tappanina selmensis 711 (Cushman), sample no. 4; 30 Globocassidulina subglobosa (Brady), sample no. 5; 712 31 Stainforthia farafraensis (LeRoy), sample no. 15; 32 Siphogenerinoides 713 eleganta (Plummer), sample no. 8; 33 Bulimina midwayensis Cushman and 714 parker, sample no. 2; 34 Bulimina farafraensis LeRoy, sample no. 10; 35715 Bulimina esnaensis LeRoy, sample no. 18; 36 Globobulimina ovata (d'Orbigny), 716 sample no. 3, 37 Coryphostoma midwayensis (Cushman), sample no. 4. 717

Plate 3. (Scale bar is 100 µm); 1 Fursenkoina sp., sample no. 2; 2 Pleurostomella 718 paleocenica Cushman, sample no. 11; 3 Stilostomella midwayensis (Cushman and 719 Todd), sample no. 4; 4,5 Valvulineria scrobiculata (Schwager), sample no. 16; 6,7 720 Eponides lotus (Schwager), sample no. 6; 8 Discorbis pseudoscopos Nakkady, 721 sample no. 7; 9,10 Cibicidoides alleni (Plummer), sample no. 4; 11,12722 Cibicidoides succedens (Brotzen), sample no. 5; 13 Allomorphina trigona Reuss, 723 sample 7; 14 Quadrimorphina allomorphinoides (Reuss), sample no. 6; 15,16724 Alabamina midwayensis Brotzen, sample no. 6; 17,18 Valvalabamina depressa 725 (Alth), sample no. 3; 19 Osangularia plummerae Brotzen, sample no. 2; 20,21 726 Oridorsalis plummerae (Cushman), sample no. 6; 22,23 Anomalinoides 727 aegyptiacus (LeRoy), sample no. 15; 24,25 Anomalinoides rubiginosus 728 (Cushman), sample no. 1; 26 Anomalinoides zitteli (LeRoy), sample no. 16; 27,28 729 Gyroidinoides subangulata (Plummer), sample no. 2; 29 Angulogavelinella 730 avnimelechi (Reiss), sample no. 11; 30,31 Gavelinella danica (Brotzen), sample 731 no. 2; 32 Karreria fallax Rzehak, sample no 4. 732

# **Table Captions**

Table 1. Benthonic foraminiferal species count from Naqb Assiut section, 734Western Desert, Egypt.735

Table 2. Habitat preferences of the abundant calcareous and agglutinated 736benthonic foraminiferal taxa at Naqb Assiut section, Western Desert, Egypt.737

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SAMPLE No.	N1	N2	N3	N4	N5	N6	N7	N8	N9	N10	N11	N12	N13	N14	N15	N16	N17	N18
Bolivinopsis spectabilis	2	4	0	2								A						
Bermudezina danica	2	0	0	0	0	3	4	2	4									
Lagena hispida	5	0	8	6	5	6	4	5	4									
Frondicularia goldfussi	4	4	2	3	2	4	3	4	3	4	3							
Karreria fallax	4	3	5	4	3	4	3	5	3	3	4							
Neoflabellina jarvisi	2	2	3	1	2	5	3	2	2	3	4	2						
Glandulina laevigata Lantiaulina	2	0	3	0	0	3	4	4	3	<b>C</b> 1	5	2						
Lenticulina pseudomamilligera	3	3	2	2	3	5	5	3	4	3	2	3						
Laevidentalina gracilis	2	0	4	5	4	5	4	3	0	4	4	3						
Marginulinopsis tuberculata	2	0	2	0	2	0	3	0	3	3	4	5						
Stilostomella stephensoni	6	5	4	0	0	3	4	4	0	2	3	3						
Ammodiscus glabratus	4	3	3	2	2	4	3	0	4	2	3	2	1					
Bathysiphon cf. discretus Ammosphaeroidina	6	3	4	4	3	6	5	6	3	4	4	3	2					
pseudopauciloculata	2	3	3	1	2	4	3	0	1	2	4	2	1					
Ramulina navarroana	7	5	6	5	0	4	5	4	0	5	6	3	0	4				
Pseudoclavulina farafraensis	5	7	6	0	5	0	3	5	6	5	2	3	5	7				
Stilostomella midwayensis	9	7	0	6	0	5	7	5	7	5	8	4	0	5	4			
Gavelinella danica	5	6	5	4	6	7	5	6	6	5	7	4	0	3	2			
Tritaxia midwayensis	3	0	7	8	6	4	6	7	8	5	6	5	4	5	6			
Clavulinoides amorpha	2	0	4	0	0	0	4	5	4	0	5	2	0	2	3			
Nodosaria (?) longiscata	3	0	5	4	0	6	0	4	5	3	4	3	1	4	5	4		
Laevidentalina colei	9	7	0	5	5	3	4	2	5	2	4	3	0	5	3	4		
Spiroloculina tenuis	6	5	3	3	4	0	3	5	2	6	11	5	0	4	6	3	5	
Loxostomoides applinae	14	21	19	16	15	17	22	19	21	19	12	4	0	3	4	4	2	

Coryphostoma midwayensis	5	0	11	8	0	0	0	5	0	7	2	2	0	5	4	3	4	
Spiroplectinella dentata	8	11	9	11	8	6	8	6	8	7	3	3	4	6	5	7	6	
Pseudoclavulina maqfiensis	4	6	0	5	4	0	0	4	5	4	1	1	0	4	6	5	3	
Cristellariopsis proinops	3	4	5	3	2	4	5	4	2	5	4	2	0	6	4	4	3	
Cibicidoides alleni	8	6	9	11	13	9	8	10	8	9	6	4	0	5	3	4	2	
Pseudonodosaria manifesta	7	4	4	0	0	0	6	0	4	5	4	3	0	5	4	5	4	
Laevidentalina cylindroides	4	0	5	0	0	3	0	0	4	2	3	1	0	3	4	7	4	
Spiroloculina proboscidea	4	6	6	3	6	7	5	6	3	3	5	2	0	4	5	9	6	
Valvalabamina depressa	4	6	9	6	7	5	7	8	7	5	6	3	0	4	4	8	7	
Valvalabamina planulata	5	4	6	4	5	3	5	5	3	3	5	2	0	3	5	7	5	
Siphogenerinoides eleganta	9	10	8	6	0	0	0	8	7	5	0	3	0	5	8	6	5	
Osangularia plummerae	8	9	7	6	6	9	5	8	6	5	3	2	0	4	4	5	6	
Lenticulina navicula	9	7	5	6	5	6	5	8	6	7	6	4	0	7	4	8	6	
Cibicidoides succedens	5	6	4	5	6	5	2	7	4	5	4	3	0	4	3	2	3	
Globocassidulina subglobosa	6	0	0	6	5	3	5	0	4	0	5	2	0	6	9	7	11	9
Alabamina midwayensis	4	6	7	5	6	8	6	5	4	9	5	3	0	2	4	3	5	4
Pyramidulina latejugata	5	4	0	6	5	7	5	3	5	4	5	3	1	6	4	3	2	4
Gyroidinoides subangulata	6	7	6	5	4	5	7	7	5	4	6	2	0	4	4	6	8	11
Gyroidinoides girardana	3	5	4	3	2	3	4	5	3	4	4	1	0	5	2	3	6	9
Valvulineria scrobiculata	3	5	2	3	5	5	6	5	4	3	3	4	0	10	15	13	14	19
Anomalinoides rubiginosus	2	1	1	1	0	0	1	1	1	1	0	1	0	0	0	1	0	1
Tappanina selmensis	7	5	6	5	0	0	0	0	0	0	0	1	0	4	4	5	8	7
Reussoolina apiculata																		
apiculata	6	0	5	4	0	3	0	4	5	3	0	3	0	4	4	3	2	5
Bulimina farafraensis	3	0	0	8	6	4	0	0	0	4	0	2	0	5	7	9	14	10
Bulimina midwayensis	7	9	8	7	10	5	7	0	3	4	0	2	0	5	6	5	7	4
Bulimina trinitatensis	4	0	0	0	5	3	0	0	0	0	6	2	0	0	0	0	0	3

Globobulimina ovata Marssonella oxycona	5	8	6	0	9	7	0	0	7	6	0	3	0	4	5	8	6	12
trinitatensis Marssonella oxycona	2	0	2	2	0	0	2	0	3	2	3	2	1	0	0	0	0	4
oxycona	4	3	4	3	6	3	4	5	4	4	3	3	2	3	2	3	0	2
Spiroplectinela esnaensis	4	4	0	6	5	0	5	4	4	3	3	2	4	5	3	5	9	8
Spiroplectinella knebeli	3	4	0	4	5	0	4	3	3	0	4	2	3	7	5	4	2	4
Dorothia bulletta	2	0	0	3	0	2	0	0	0	3	5	4	2	3	4	3	2	2
Gaudryina textulariformis Quadrimorphina	4	0	0	0	6	4	5	3	6	5	5	3	0	4	4	8	6	4
allomorphinoides	3	4	6	0	0	0	0	0	4	4	0	2	0	3	3	5	3	4
Saracenaria triangularis	3	3	0	0	3	0	4	3	0	4	9	4	0	2	2	1	3	3
Lenticuina midwayensis	6	5	7	8	9	6	9	8	7	10	5	3	0	3	6	5	7	8
Lenticulina turbinata	4	4	5	7	5	6	5	6	5	6	6	4	0	3	5	7	3	4
Anomalinoides zitteli	2	3	2	4	7	0	0	4	3	3	2	1	0	9	14	13	10	8
Anomalinoides spissiformis	3	4	7	5	4	3	6	5	4	7	4	3	0	5	4	3	2	3
Anomalinoides aegyptiacus	2	3	3	7	6	2	0	1	2	3	2	2	0	9	15	11	10	13
Oridorsalis plummerae Angulogavelinella	5	4	5	7	4	8	5	6	5	4	6	3	0	6	8	7	5	9
avnimelechi		2	1	1	Ι	0	1	0	1	2	3	1						
Neoflabellina paleocenica		4	5	2	4	4	1	3	3	2	2	2						
Discorbis pseudoscopos		3	5	4	6	4	5	7	5	4	5	3	_					
Ammoglobigerina altiformis		2	1	2	1	3	1	2	1	0	4	2	3					
Pyramidulina paupercula		4	5	0	4	6	4	4	5	4	4	2	1	4				
Gaudryina elegantissima		6	0	5	4	3	4	4	5	0	0	3	5	6	2			
Fursenkoina sp.		5	0	0	4	6	0	0	0	0	0	2	0	6	8	6	7	13
Bulimina esnaensis		2	0	4	0	0	0	2	0	0	0	2	0	3	4	5	8	11
Pyramidulina vertebralis		3	7	5	4	4	3	5	6	7	5	3	1	3	4	3	2	4

Bulimina quadrata		4	0	3	0	4	0	0	4	5	3	3	0	4	3	5	4	6	
Vulvulina colei		5	0	0	0	0	0	0	3	3	0	2	2	2	3	0	0	4	
Anomalinoides umbonifera		2	3	2	2	2	3	1	1	2	2	1	0	5	3	7	10	15	
Karrerulina conversa			1	0	0	0	0	0	0	0	0	0	2	2					
Allomorphina trigona				3	0	3	4	0	0	0	0	0	0	2	0	3	2	3	
Trifarina esnaensis				5	10	0	5	0	0	0	0	0	0	3	5	8	12	11	
Clavulinoides asper whitei					4	3	2	0	0	0	0	1							
Gaudryina pyramidata					4	3	2	0	0	0	3	2							
Eponides lunatus					2	3	1	4	2	2	4	0	0	4	5	5	6	9	
Eponides lotus					2	4	3	2	1	2	4	2	0	3	4	7	8	11	
<i>Cibicidoides pseudoacutus</i>						6	8	6	3	7	6	4	0	2	2				
Clavulinoides trilatera								3	4	2	2	1	0	3	3				
Stainforthia farafraensis								$\langle \rangle$			2	3	0	11	16	13	14	19	
Pleurostomella paleocenica											2	0	0	2	0	3	5	2	
Eouvigerina aegyptiaca								Y						5	3	2	8	9	
Total number of						Â													
individuals	300	300	300	300	300	300	300	300	300	300	299	202	45	294	300	303	302	291	
			A	Ĉ	3	1													

## **Epifaunal calcareous**

<u>Rounded trochospiral</u> Anomalinoides rubiginosus <sup>1</sup> Gavelinella danica

<u>Plano-convex trochospira</u>l Angulogavelinella avnimelechi Anomalinoides zitteli Cibicidoides spp.<sup>1</sup> Gyroidinoides girardana<sup>1</sup> Valvulineria scrobiculata<sup>2, 4</sup> Gyroidinoides subangulata Valvalabamina spp. Alabamina midwayensis

Biconvex trochospiral Anomalinoides aegyptiacus <sup>3, 8</sup> A. spissiformis A. umbonifera Cibicidoides pseudoacutus Eponides lunatus Eponides spp. Cibicidoides alleni C. succedens Oridorsalis plummerae <sup>4</sup> Osangularia plummerae

Biconvex planispiral Lenticulina spp. <sup>1, 6, 7</sup>

Discorbis pseudoscopos

<u>Concavo-convex trochospiral</u> Karreria fallax

<u>Milioline</u> Quinqueloculina Spiroloculina spp.

<u>Palmate</u> Frondicularia spp. Neoflabellina spp.

**Epifaunal agglutinated** <u>Tubular or branched</u> Bathysiphon <sup>5, 8</sup>

 $\frac{\text{Coiled flattened} - \text{streptospiral}}{\text{Ammodiscus spp.}^{3}}$ 

Rounded trochospiral Ammosphaeroidina sp. ?

# Infaunal calcareous

Cylindrical tapered Bulimina spp. Fursenkoina sp. Globobulimina ovata Glandulina laevigata? Laevidentalinids **Pyramidulinids** Nodosariids Pseudonodosaria manifesta Siphogenerinoides eleganta Eouvigerina aegyptiaca Pleurostomella spp. Stilostomella sp. Trifarina esnaensis Stainforthia farafraensis Tappanina selmensis

<u>Coiled-rectilinear</u> Saracenaria spp. Marginulinopsis sp. Cristellariopsis proinops

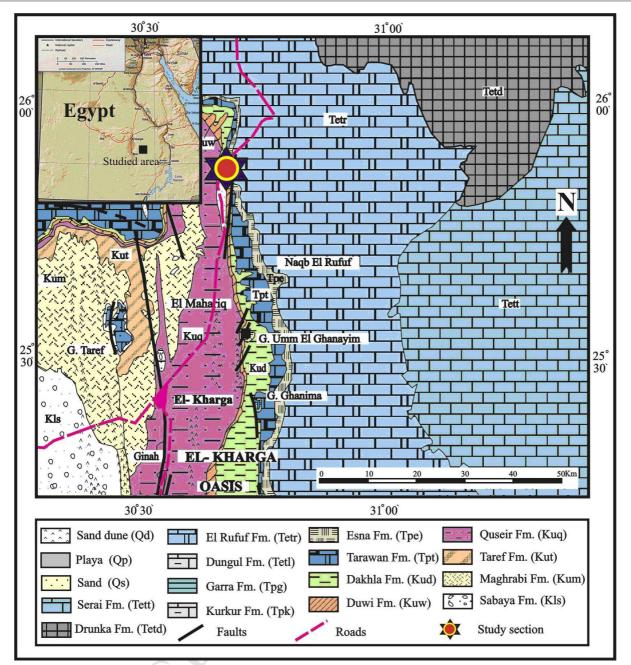
<u>Flattened tapered</u> Loxostomoides applinae Coryphostoma midwayensis

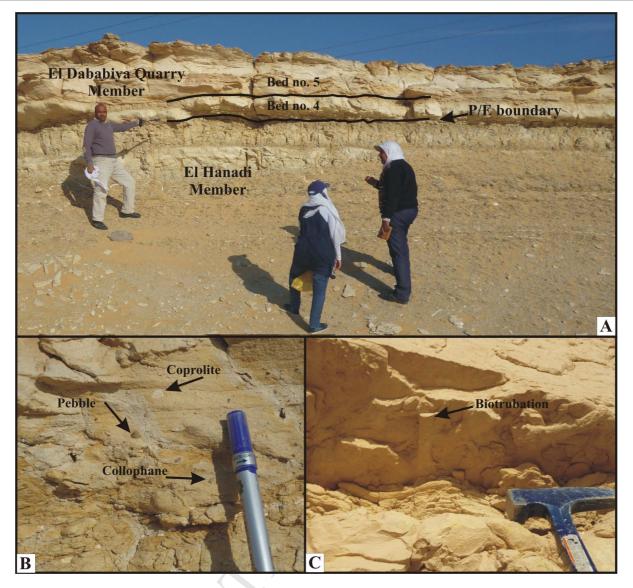
<u>Spherical – globose</u> Lagena spp.<sup>1</sup> Globocassidulina subglobosa morphina allomorphinoides Allomorphina sp. Reussoolina spp.<sup>1</sup>

<u>Irregular</u> Ramulina navarroana

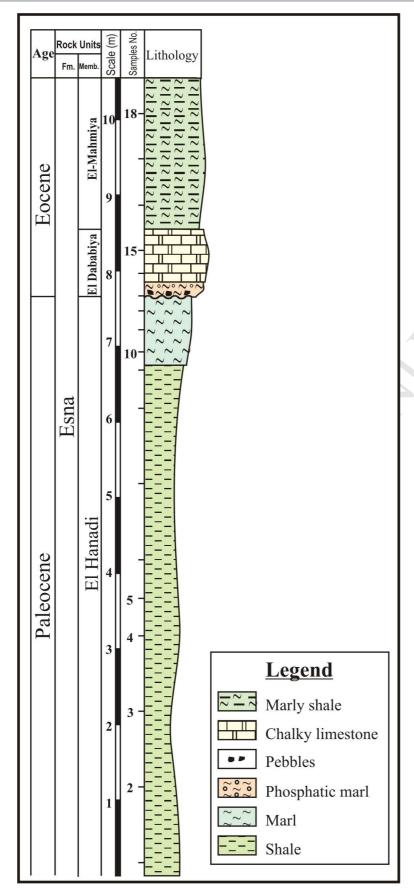
Infaunal agglutinated <u>Elongate multilocular</u> *Gaudryina* spp. *Bolivinopsis spectabilis Vulvulina* spp. *Karrerulina conversa Spiroplectinella dentate*  Spiroplectinella esnaensis S. knebeli Clavulinoides amorpha C. trilateral Dorotia bulleta Gaudryina pyramidata<sup>8</sup> Marssonella oxycona<sup>8</sup> Pseudoclavulina spp. Tritaxia midwayensis Bermudezina danica

<sup>1</sup>Widmark and Malmgren <sup>2</sup>Speijer (1992); and Schmitz (1998); <sup>3</sup>Speijer and Wagner (2002); <sup>4</sup>Widmark and Speijer <sup>5</sup>Peryt (1997); et al. (1997); <sup>6</sup>Corliss (1991); <sup>7</sup>Thomas and Schackleton (1996); and <sup>8</sup>Kaminski et al. (1996).



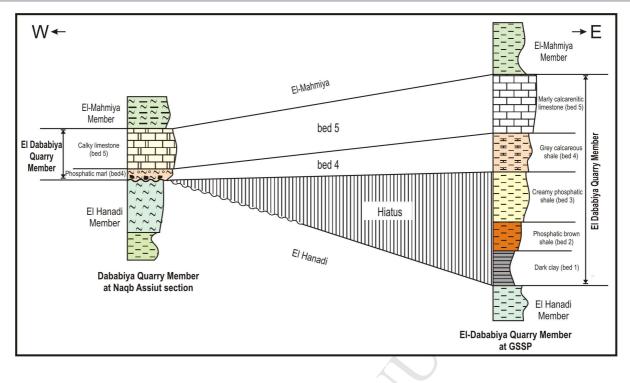






Age		Present work qb Assiut		Abdel- Razik (1972)	Hermina 1990 North Kharga		ipuis <i>et al.</i> (2003) Dababiya section		ubry <i>et al</i> . (2007) Dababiya section
		Not studied	Thebes Fm.	Abu Had Mb.			Esna Unit 3		Abu Had Mb.
Eocene	Esna Fm.	El-Mahmiya Mb. Dababiya Bed 5 Bed 4 Quarry Mb.	Shale	El-Shaghab Mb.	Esna Fm.	Esna Shale	Estimation Control Formation C		El-Mahmiya Mb. Dababiya Bed 5 Bed 4 Quarry Bed 3 Bed 2 Bed 2 Bed 1
aleocene		El Hanadi Mb.	[ r_]	El Hanadi Mb.			Esna Unit 1		El Hanadi Mb.
Palec	Not studied			rawan Chalk	Tarawan Fm.		Tarawan imestone	r -	Farawan Fm.

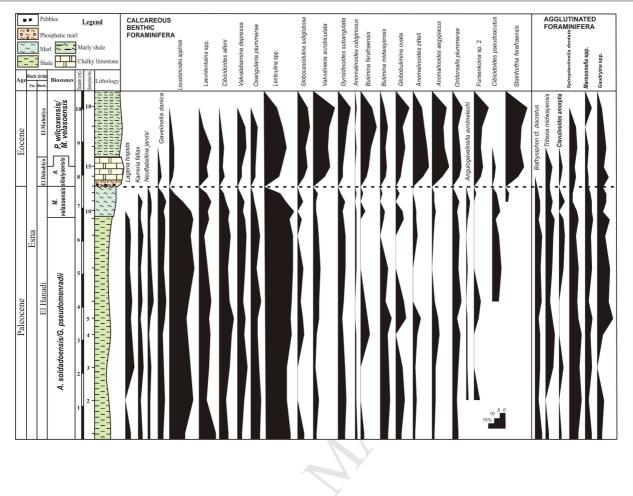
arawan Tara \_\_\_\_\_\_\_ Tara \_\_\_\_\_\_ Lime

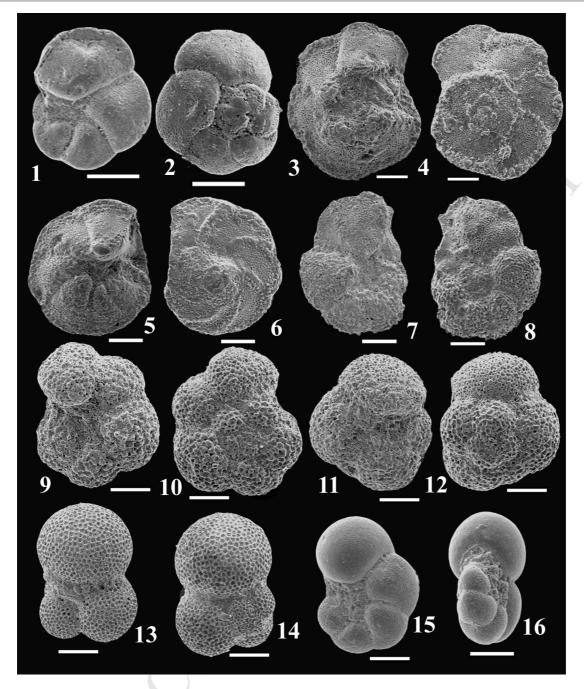


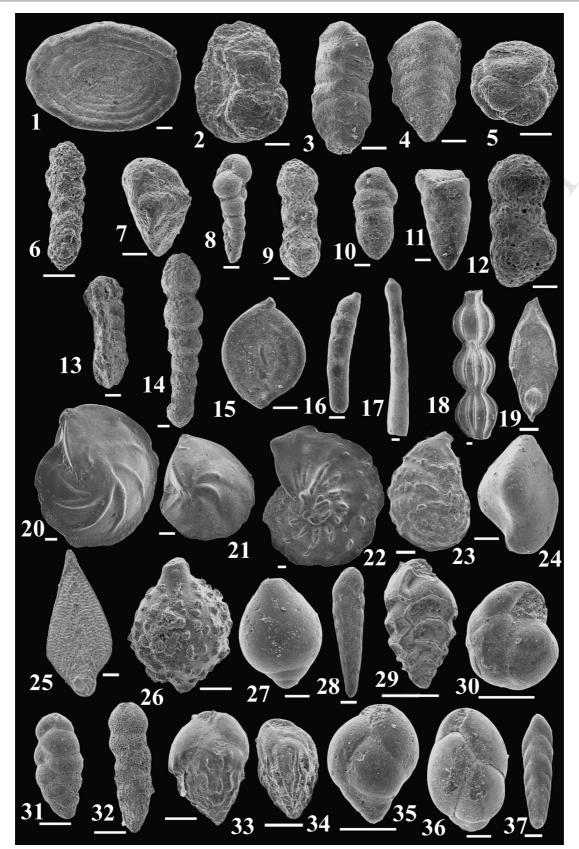
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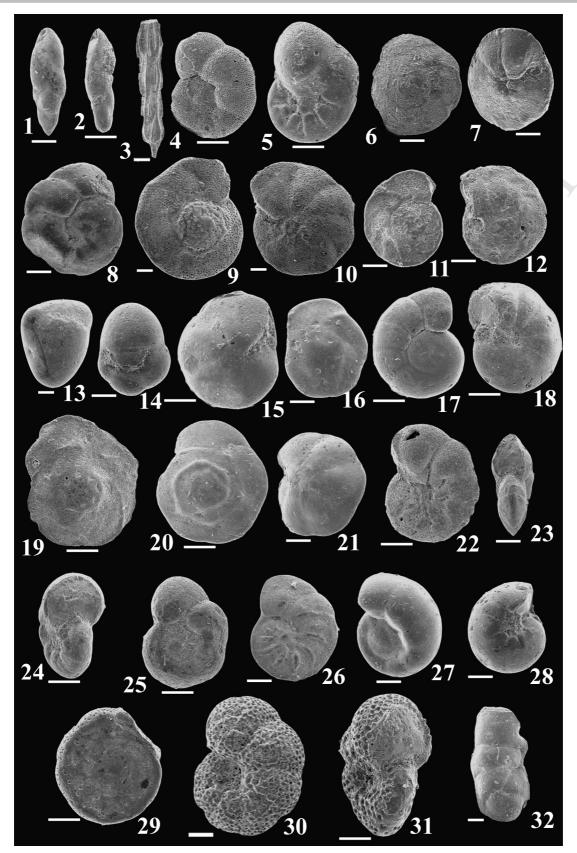
	Eocene
	Esna
A. soldadorensisG. perudomentali a. " a. " a. " a. " b.	El Dababiya El-Mahmiya 🐺
	M. A. D. wilcovoncis/
	velasoensis sibaiyaensis M. velasoensis
	1
	88 88 6 6 4 4 3 3
	<u>التي</u> 118- - - - - - - - - - - - - - - - - -
	M.anoulata
	Mupunesmu
	M.conicotruncata
	<i>M.gracilis</i>
	Monstorms
Pseudolustigeritu wilcovenis	
Pseudolustigeritu wilcovenis	- M.subbotina
Pseudolastigerita wilcovensis	M.velascoensis
	A africana
	A.coalingensis
	A.esnaensis
	——————————————————————————————————————
	4. mckannai
	A.SibaylaenSis
	A.Strabocella
	A trinlex
	Scancellata
	Seocaena
	C inconient
	Strangulars
	S. triloculinoides
	S trivalis
	C unfreenence
	Parasubbotina varianta
	I.beudermana
	I.lodoensis
	G.imitata
	Geovalis Geovalis
	Gulanoconica
Pseudohastigerina wilcoxensis	
	Pseudohastigerina wilcoxensis

Age	Datum events	The present study	Ве (19	erggren <i>et al.</i> 995) Standard	(2003	gren & Ouda 3) Dababiya on (GSSP)		Berggren & earson (2005)	,	Wade <i>et al</i> . (2011)
ne	M. velascoensis*	P. wilcoxensis- M. velascoensis				P5c	E2	P. wilcoxensis- M. velascoensis	E2	P. wilcoxensis- M. velascoensis
Eocene	P. wilcoxensis ← A. sibaiyaensi &	A. sibaiyaensis	Р5	M. velascoensis	Р5	P5b	E1	A. sibaiyaensis	E1	A. sibaiyaensis
cene	A. africana 🛀 G.pseudomenardii	M. velascoensis				P5a	P5	M. velascoensis	P5	M. velascoensis
Paleocene	A. soldadoensis	A. soldadoensis- G.pseudomenardii	P4c	A. soldadoensis- G.pseudomenardii	P4	P4c	P4c	A. soldadoensis- G.pseudomenardii	P4c	A. soldadoensis- G. pseudomenardii
							S			









-The lithostratigraphy of the P-E transition at Naqb Assiut, Western Desert, Egypt is studied.

-The biostratigraphy of this transition is carried out and four planktonic foraminiferal zones are defined.

-The P/E boundary is documented and discussed.

-The benthonic foraminiferal fauna during this interval are analyzed and discussed.

-The Paleobathymetrical and paleoenvironmental changes across the P/E boundary are explained.

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